Snow Hydrology and Modelling in Alpine, Arctic and Forested Basins



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Mountain Snow

CMC 150 ¥.> 100 75 50 40 30 20 10 5 Janvier 0 January

Snow depth in January



Snow depth in June

vast water reserves in winter snowpack

Study Elements

Processes

- Snow accumulation, structure and observation
- Turbulent transfer to snow
- Radiation effects on snowmelt under tundra shrubs and evergreen forests

Parameterisations

- Blowing snow over complex terrain
- Irradiance in complex terrain longwave from terrain, shortwave shadows
- Forest snow interception, unloading and sublimation
- Sub-canopy snowmelt
- SCA Depletion in complex terrain,
- Contributing area for runoff generation in snowmelt period
- Prediction
 - Wind and atmospheric modelling over complex terrain
 - Level of spatial complexity necessary in models
 - Regionalisation of CLASS parameters
 - Snow modelling contribution to MESH
 - CRHM
 - Arctic and sub-arctic snow hydrology, Wolf Creek & Trail Valley Creek
 - Alpine snow hydrology, Marmot Creek
 - Montane forest snow hydrology, Marmot Creek

Blowing Snow in Complex Terrain



Inter-basin water transfer

Transport of snow to drifts

Supports glaciers, late lying snowfields, hydrological contributing areas



Computer simulation of wind flow over mountains



Granger Basin, Wolf Creek, Yukon

Simulation of Hillslope Snowdrift



3 km

Marmot Creek Research Basin



Kananaskis River valley





CRHM Mountain Structure



Alpine Hydrological Response Units



Winter Snow Redistribution Modelling



Winter Snow Redistribution and Sublimation





Point Evaluation of Snowmelt Model



Frequency Distributions of SWE from LiDAR Depths and Measured Density



SWE distribution within HRU fit log-normal density distribution

Snowcovered Area from Oblique Terrestrial Photographs, Aerial Photographs and LiDAR DEM



Snow-covered Area Depletion Modelling

Four HRU (NF, SF, EF, VB) with modelled melt applied to SWE frequency distributions.



Observed – using oblique photography

Uniform – spatially uniform SWE distributions and applied melt rates for each HRU Variable SWE dist. – each HRU has a distinct distribution of SWE Variable snowmelt – each HRU has a distinct melt rate applied Fully distributed – each HRU has a distinct distribution of SWE and applied melt rate



Visualisation of Snowmelt Runoff Intensity

Early Snowmelt Period - 2008







Snow Interception & Sublimation



Net Radiation to Forests: Slope Effects



Forest Snow Regime on Slopes



Open slopes highly sensitive to irradiation difference, forests are not







HRU Delineation

- Driving meteorology: temperature, humidty, wind speed, snowfall, rainfall, radiation
- Blowing snow, intercepted snow
- Snowmelt and evapotranspiration
- Infiltration & groundwater
- Stream network

Model Structure



Model Tests - SWE



Snow Accumulation at South-facing Bottom Slope of Fisera Ridge, Marmot Creek



Snow Accumulation at Ridgetop of Fisera Ridge, Marmot Creek



Streamflow Prediction 2006







Mean Bias = -0.13 all parameters estimated from basin data

Streamflow Prediction 2007

Marmot Creek Daily Discharge





Mean Bias = -0.068 all parameters estimated from basin data

Conclusions

- Appropriate process based models driven by enhanced remote sensing and good observations can be used to achieve adequate hydrological prediction in the alpine.
- Model process and spatial structure must be appropriate to the complexity of the energy and mass exchange processes as they operate on the landscape.
- It is possible to test for the most appropriate structure for balance between model complexity and predictive ability.