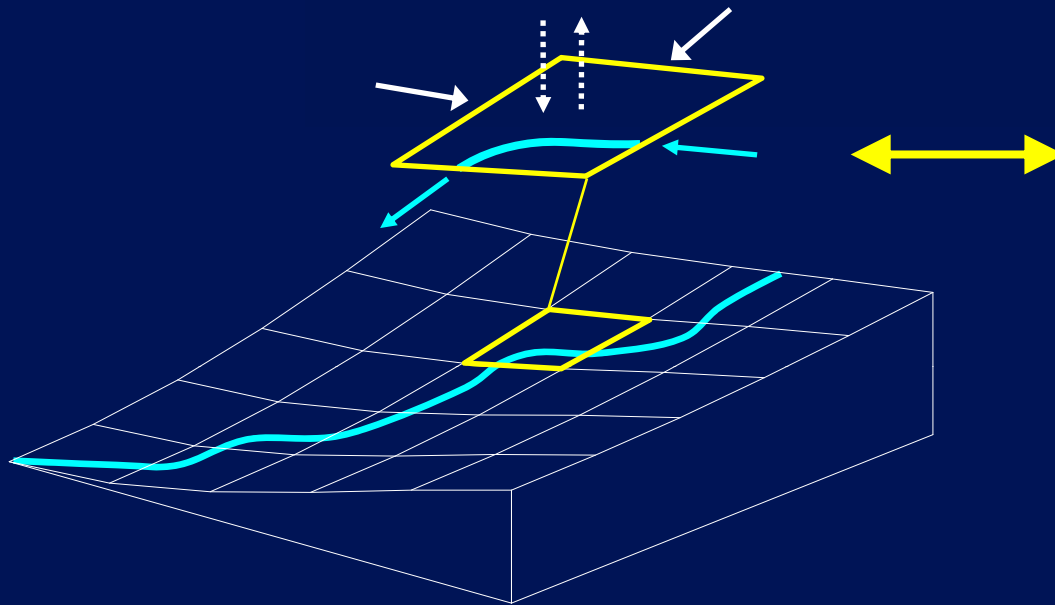


What is Hydrological Parameterization?



Scotty Creek Basin

Abstraction of complex processes into a model.
→ simplification, scaling up, “fitting” (fuzzy!)

The model must capture the **ESSENSIAL FEATURE** of physical processes.

How do we identify the Essential Feature?

- Field "intelligence" gathering
- Knowing what to look for
 - Hydrological process
 - Equations and models
- Look at big and small picture



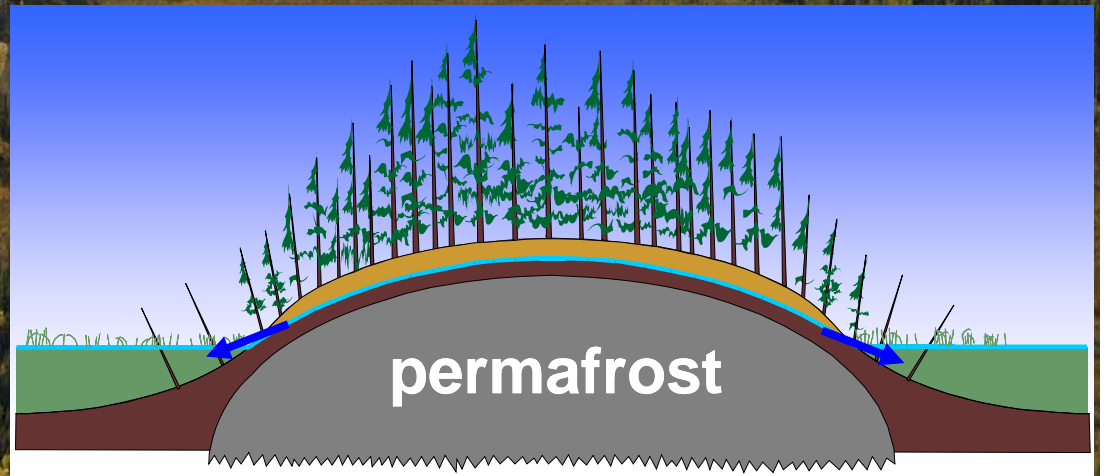
This is the best (and most fun) part of hydrological science!

Hydrologically Distinct Land-Cover Types

flat bog

peat plateau

channel fen



Channel Fen Parameterization

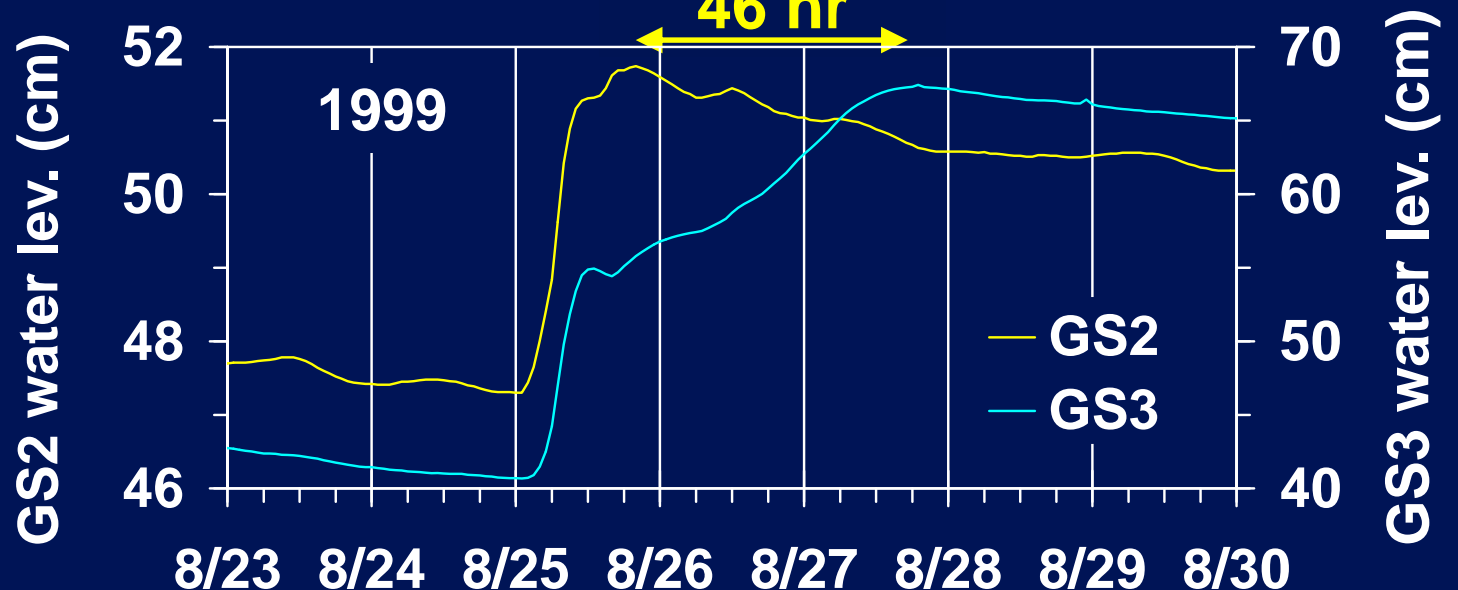
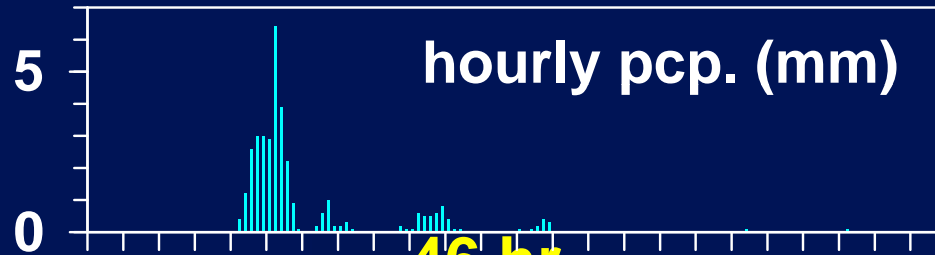
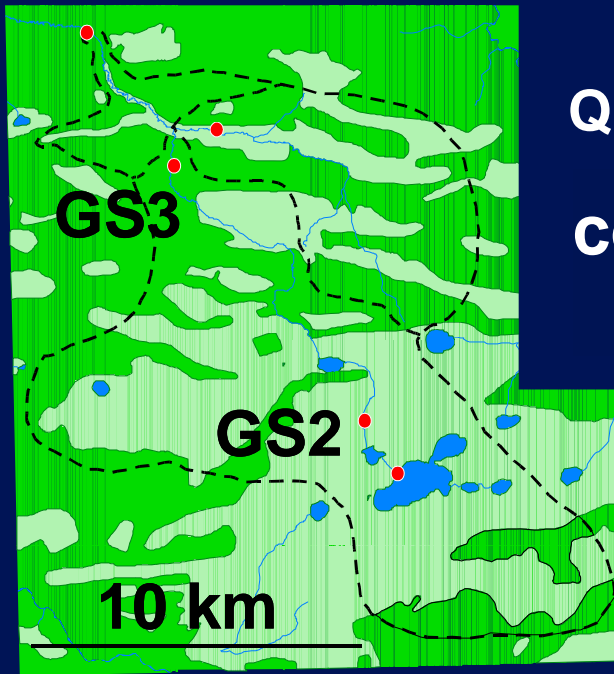
Quinton *et al* (2003. Hydrol. Proces. 17:3665)

$$\text{celerity} = 1.7 \times (\text{depth})^{2/3} \times (\text{slope})^{1/2} / n$$

↑ wave propagation

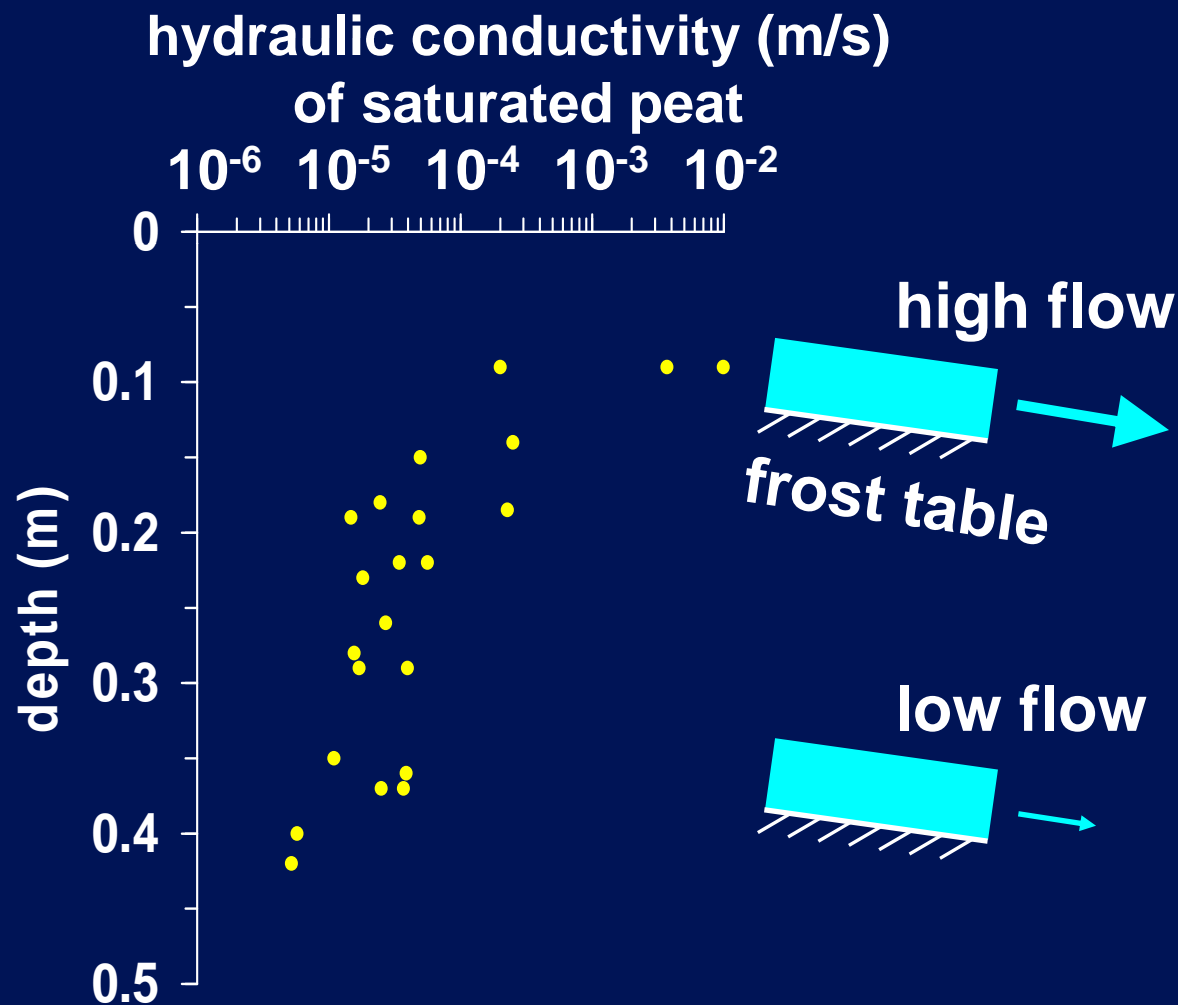
Manning ↑

$$n = 0.13-0.17$$

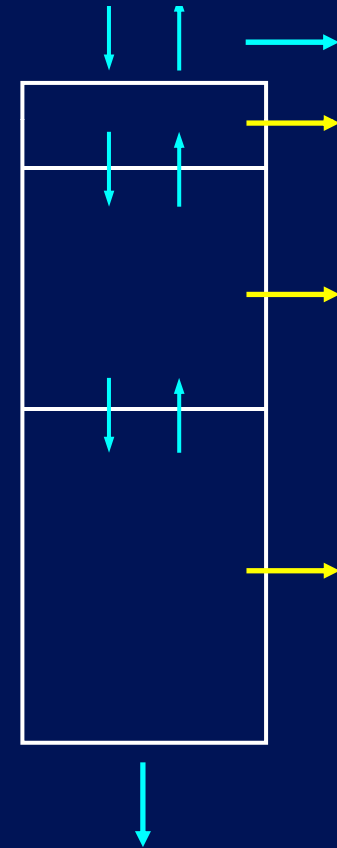


Peat Plateau Runoff Parameterization

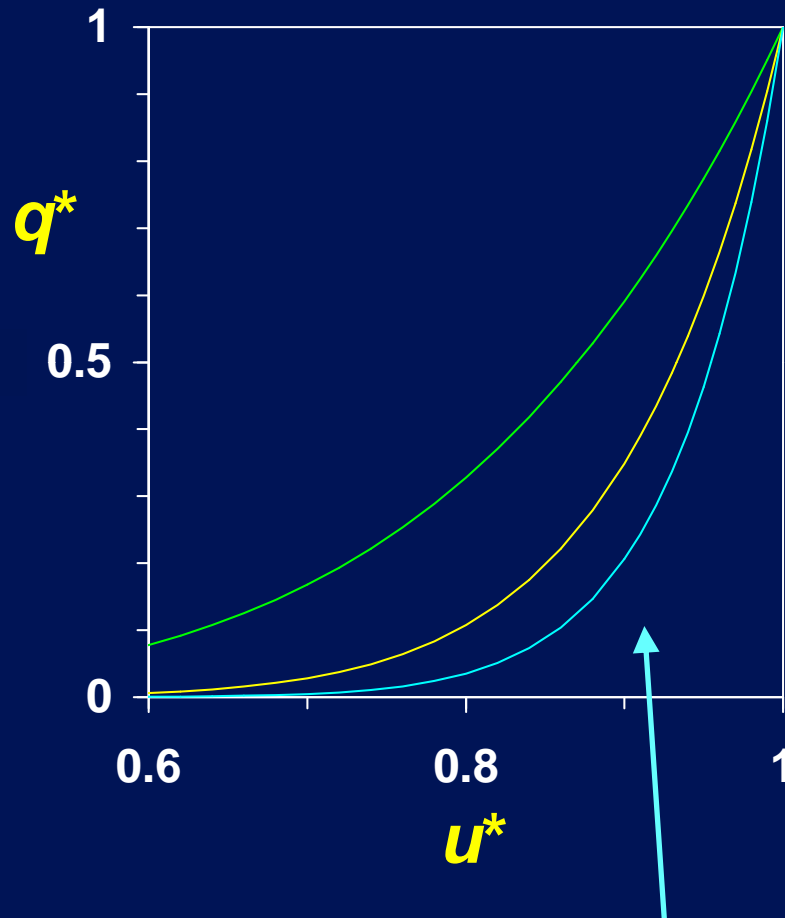
- Majority is subsurface, above the frost table
- Direction is lateral, not vertical drainage



CLASS3.1
(Soulis et al., 2000)



Lateral Drainage in CLASS 3.1



Drainage flux per area, q

Maximum drainage, q_{max}

→ Complete saturation.

Depends on slope angle,
hydraulic conductivity.

Average water storage, u

Maximum storage, u_{max}

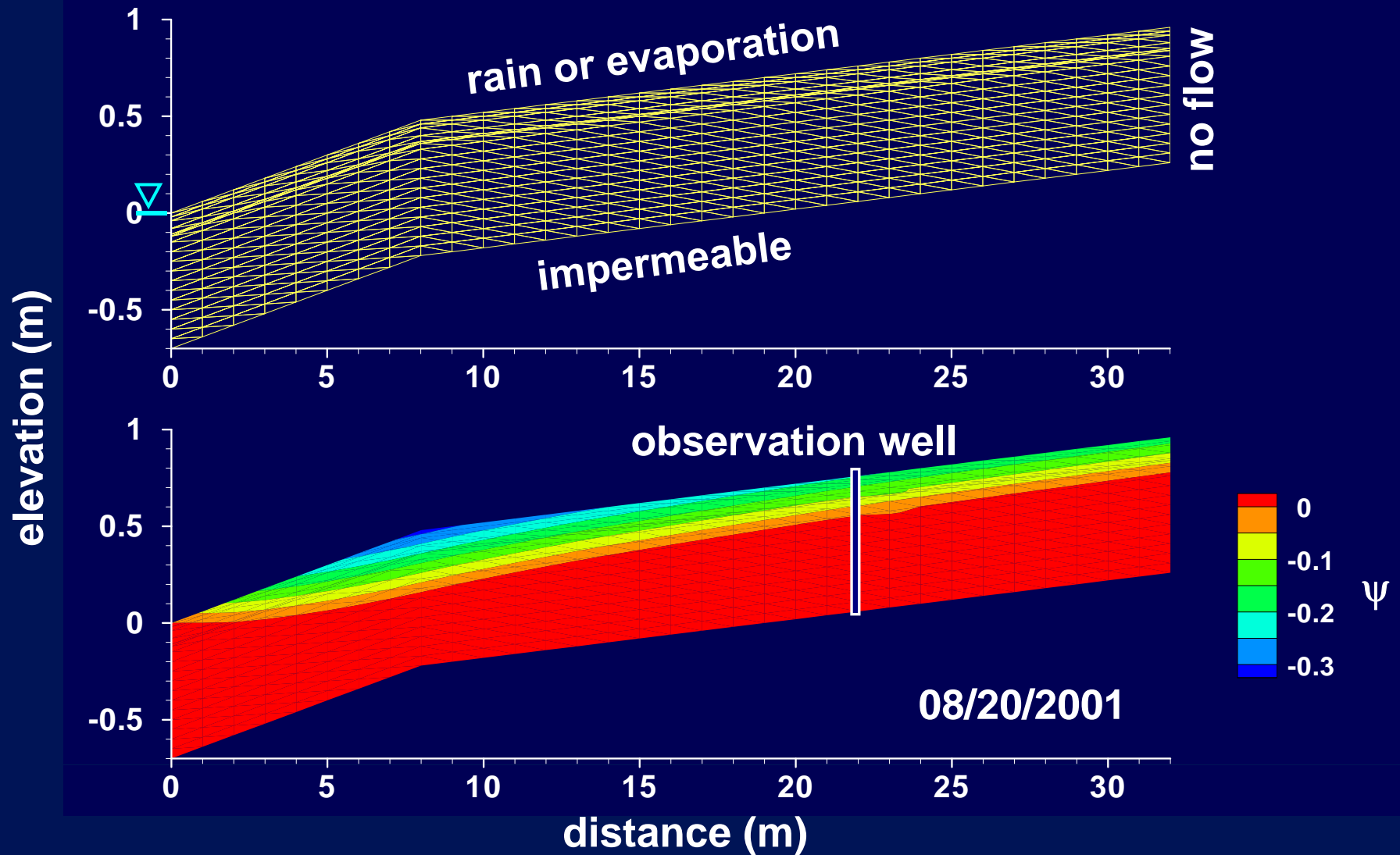
→ Complete saturation.

Normalized drainage $q^* = q/q_{max}$

Normalized storage $u^* = u/u_{max}$

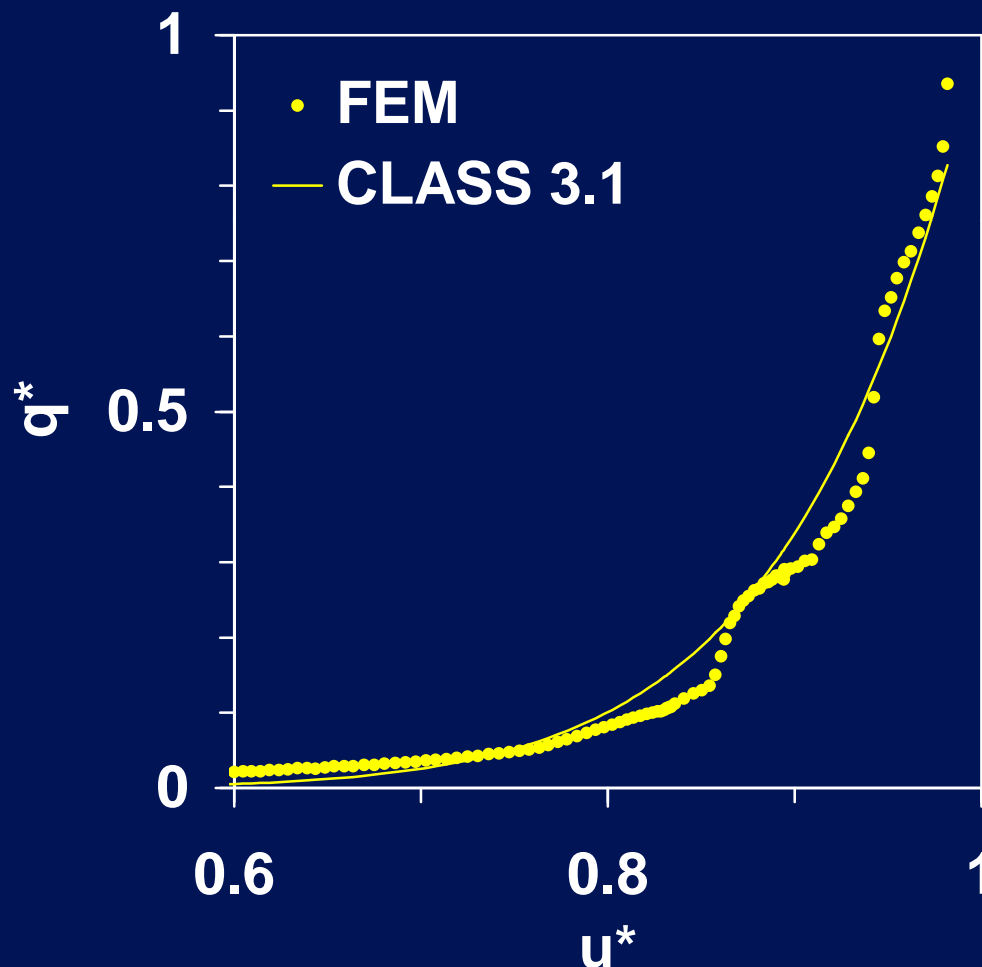
Complex interplay
of many processes

Finite Element Variably Saturated Flow Model Princeton UNSAT-2D



Finite Element Model (FEM) as a Virtual Slope

- Verify the detailed slope model against field data.
- Then, use it to parameterize a basin model.



CLASS 3.1

Analytically derived q^*-u^* .

FEM

1. Numerical drainage experiment.
2. Derive numerical q^*-u^* .
3. Determine equivalent q^*-u^* for CLASS 3.1.

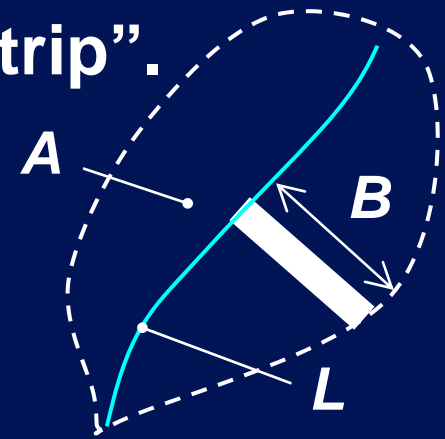
Marmot Creek Baseflow Parameterization

Analytical Solution of Brutsaert (2005. Hydrology)

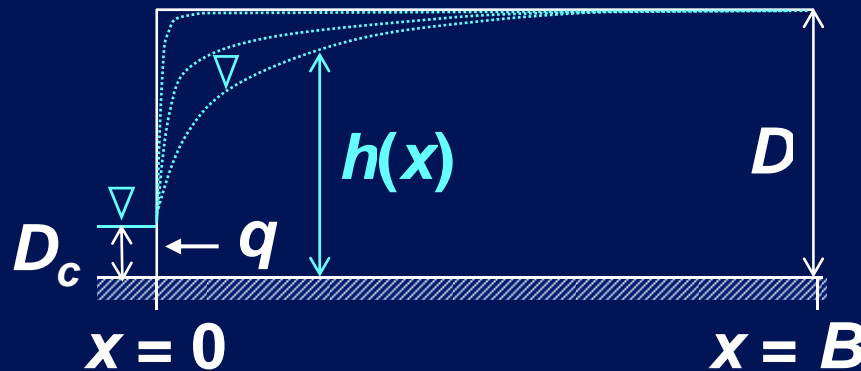
Boussinesq equation for a hillslope “strip”.

$$\frac{\partial}{\partial x} \left(Kh \frac{\partial h}{\partial x} \right) = S_y \frac{\partial h}{\partial t}$$

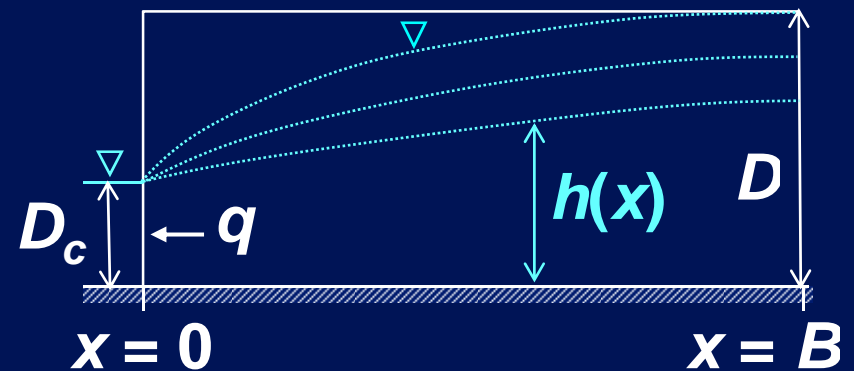
K : conductivity
 S_y : specific yield



early time solution



late time solution

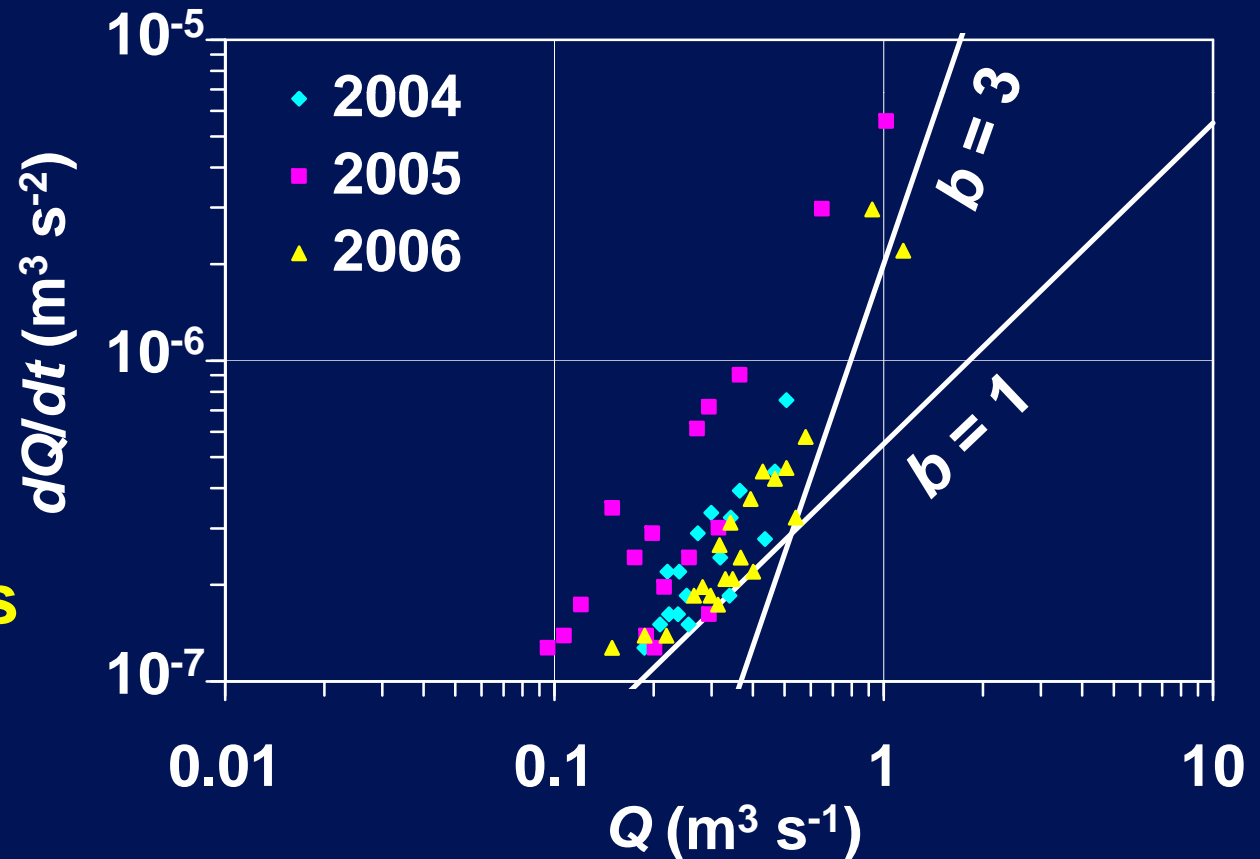


q : flow per shore length ($\text{m}^2 \text{s}^{-1}$)

$$\frac{dQ}{dt} = -aQ^b$$

Time derivative of baseflow discharge (Q) is proportional to Q^b .
→ $b = 3$ for early time, 1 for late time.

K can be determined from the intercepts of two envelope curves.



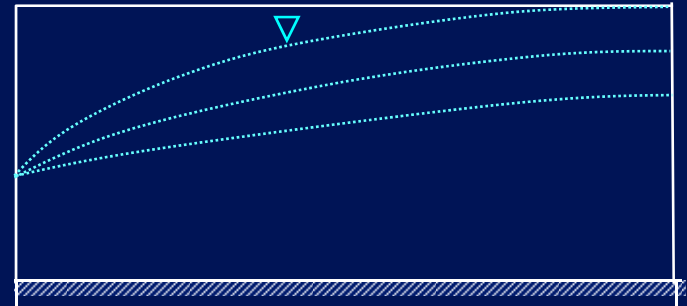
$K = 3 \times 10^{-3} \text{ m/s}$

Reasonable??

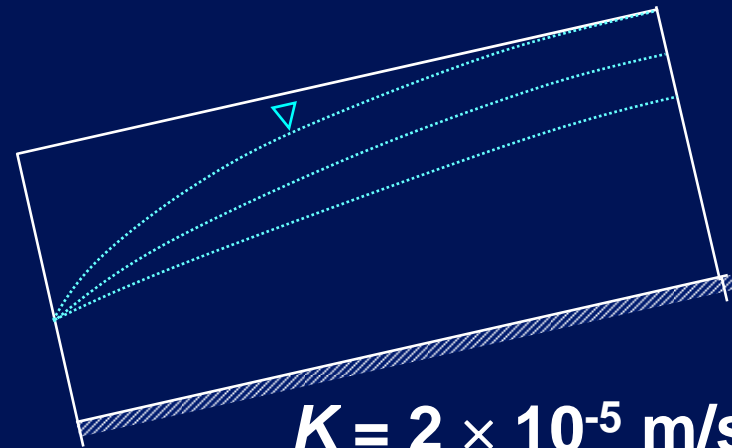
Scale-Dependent Conductivity?



photo by J. Pomeroy



$$K = 3 \times 10^{-3} \text{ m/s}$$



$$K = 2 \times 10^{-5} \text{ m/s}$$

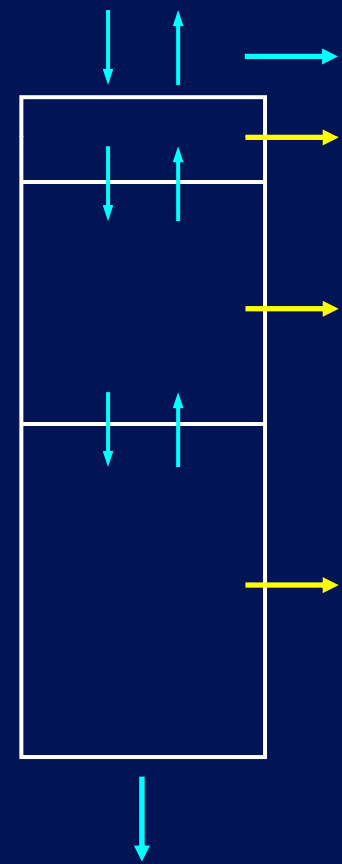
No, it is “model-dependent” conductivity!

Lake O'Hara Research Basin

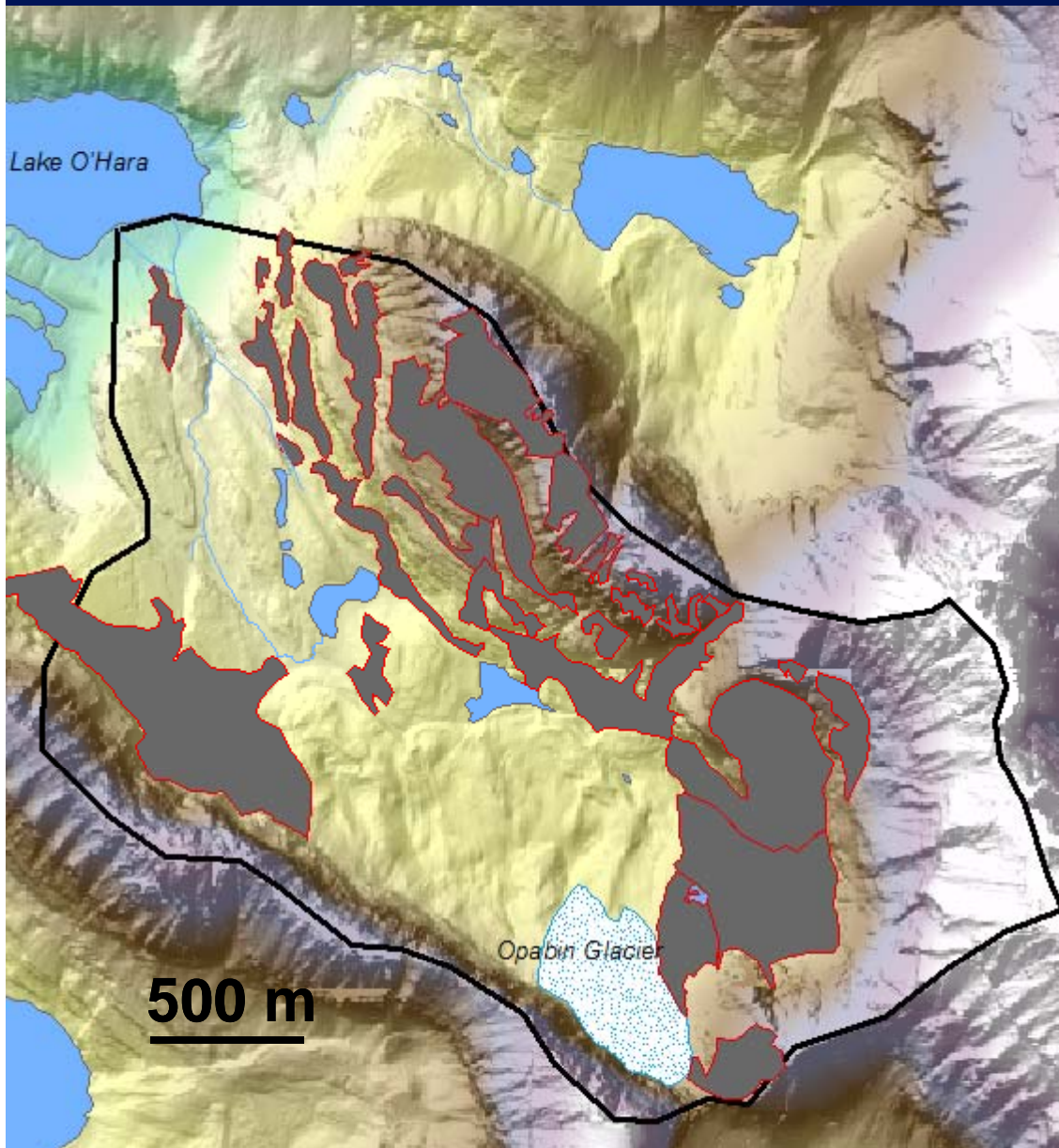
Parameterization? What? How?



MESH



Major Hydrological Land-Cover Types



**Example:
Talus slopes**

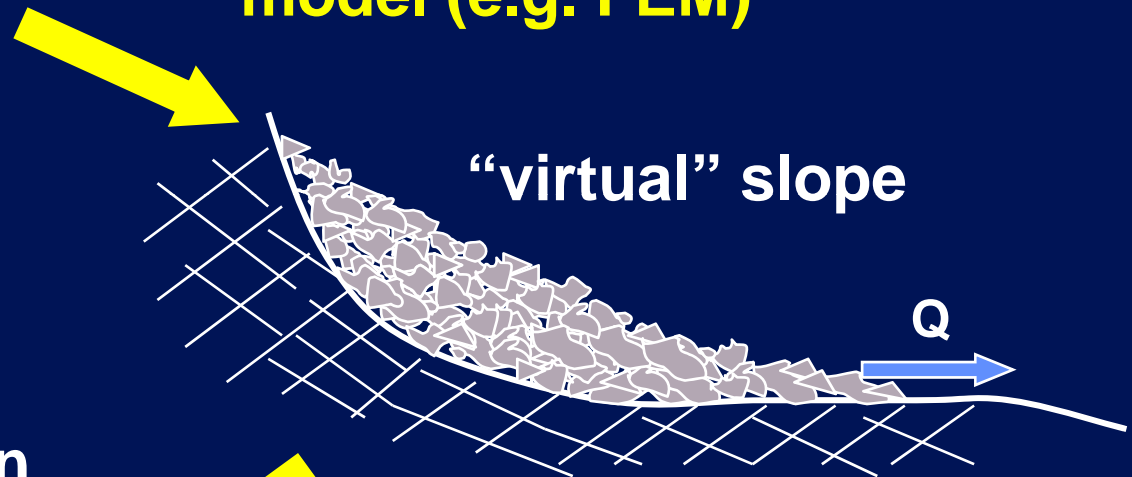


From Processes to Parameterization

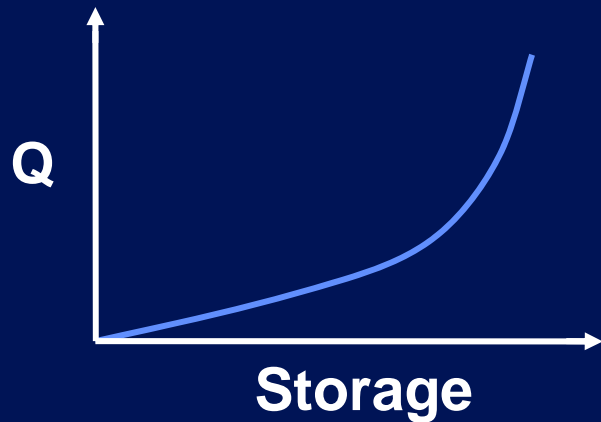
Field observation



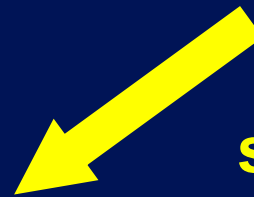
physically-based
model (e.g. FEM)



grid-scale function



simulation and
sensitivity analysis



Take Home Messages

1. Parameterization by a detailed, **“virtual” model**.
2. Scale-dependent vs. **model-dependent** parameter.
3. Process people need to understand the equations and models **VERY WELL**.
4. Modellers need to make the algorithm **transparent** to process people:
 - Consistent with published papers.
 - No arbitrary “tricks”** in the code.