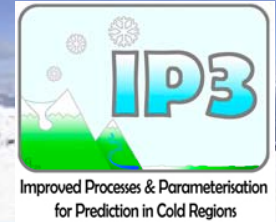




Canadian Foundation for Climate  
and Atmospheric Sciences (CFCAS)  
Fondation canadienne pour les sciences  
du climat et de l'atmosphère (FCSCA)



## Exploring Hydrologic Similarity in the Marmot Creek Basin



<http://www.geography.ryerson.ca/wayne/thesis.htm>



By:

Ken Snelgrove and Yanhzen Ou  
Memorial University of Newfoundland

For:

IP3 Third Annual Workshop  
Whitehorse, Yukon  
November 12-15, 2008



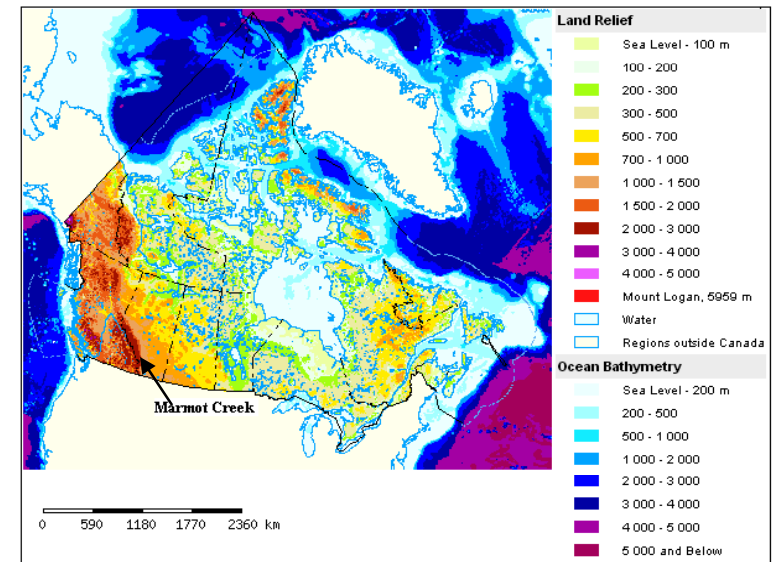


## Outline

- ◆ Marmot Creek
- ◆ Review of Last Years Presentation
- ◆ New Data Sources and the need for Hydrologic Similarity
- ◆ Topographic Index Calculation
- ◆ Evaporative Resistance Simulations using CLASS
- ◆ Potential Evapotranspiration Estimation using Penman-Monteith
- ◆ TOPMODEL Simulation and Genetic Algorithm Calibration
- ◆ Relations between Static Similarity Measures  
land cover vs. topographic index
- ◆ Future Work



# Marmot Creek Basin

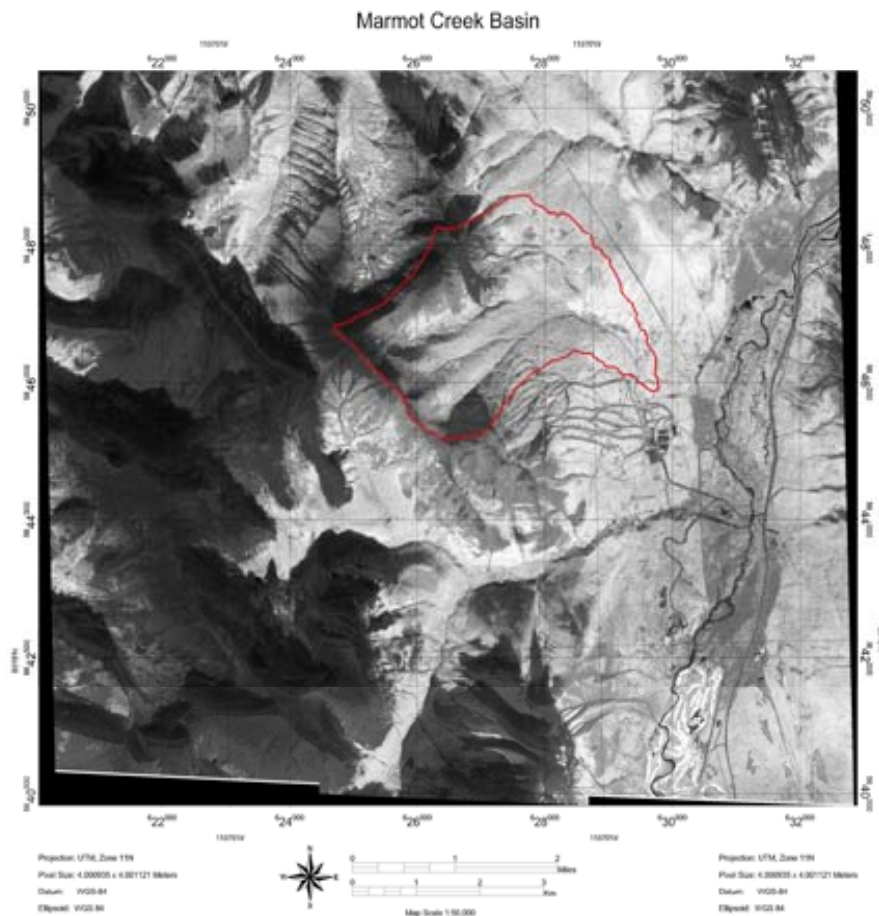


- ◆ Located at latitude 50°57'N and longitude 115°10'W
- ◆ About 110km southwest of Calgary



## Marmot Creek Basin

- ◆ Total area : 9.5km<sup>2</sup>
- ◆ Elevation: 1585-2805m, the mean value is 2112m, the difference of elevation is 1220m





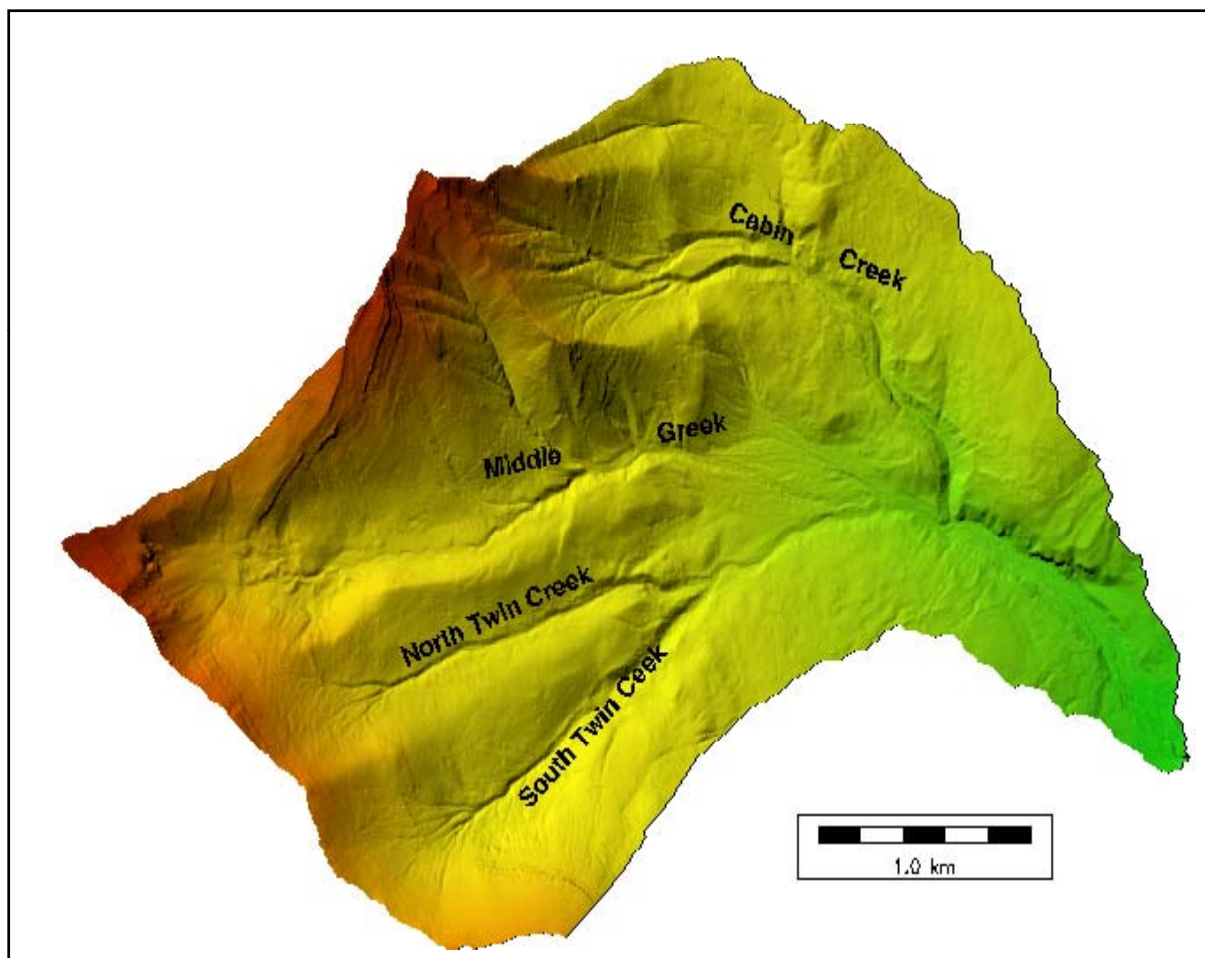
## Marmot Creek Basin

- ◆ **Three sub-basins:**

Twin Creek

Middle Creek

Cabin Creek





## Marmot Creek Basin

- ◆ **Forest Cover: 60%**
- ◆ **Soil types:** Brunisolic Grey Wooded Soils, podzolic soils, regosolic soils, alpine black soils, local gleysolic and organic soils (Stevenson, 1967(thesis))
- ◆ **Climate Characteristic:**
  - Mean annual precipitation—1080mm
  - Average July temperature— 2 to 18°C
  - Average January temperature — -6 to -18°C
- ◆ **Streamflow Characteristic:**
  - Groundwater is the main source
  - In midsummer, 70% derived from snow melting
  - Mean Annual Runoff = 425 mm



## Data Collection

- ◆ **Topographic Data**

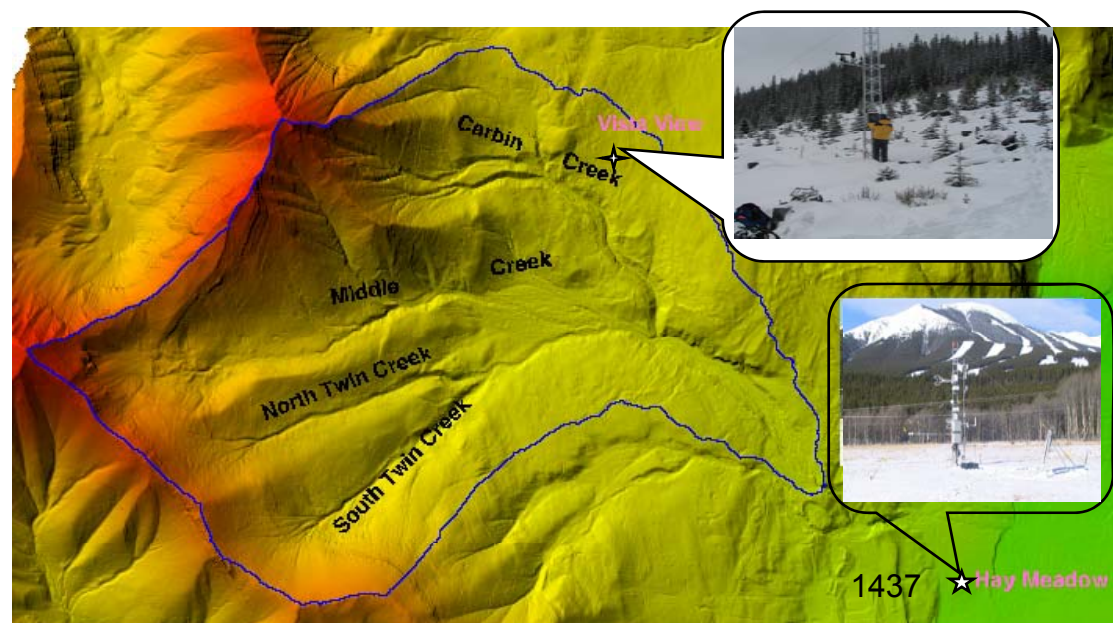
Shuttle Radar Topography Mission (SRTM) data – 90m

Light Detection and Ranging (LiDAR) data – 1m

- ◆ **Meteorological Data**

Hay Meadow Station

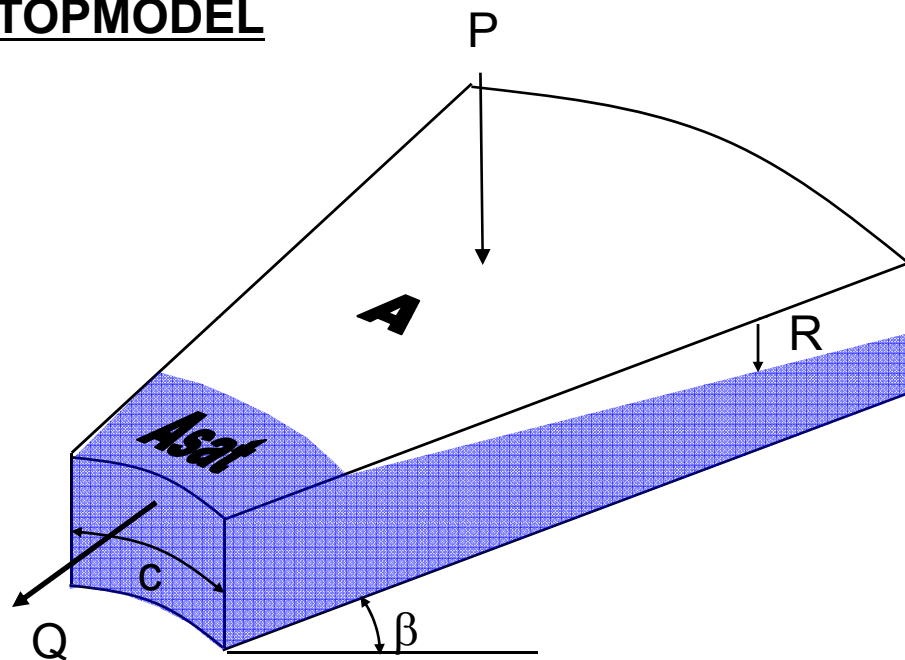
Vista View Station





# Last Year in Review

## TOPMODEL



Continuity:  $I - O = \frac{dS}{dt} = 0$  (Steady Assumption)

$$AR - Q = 0$$

$$Q = AR$$

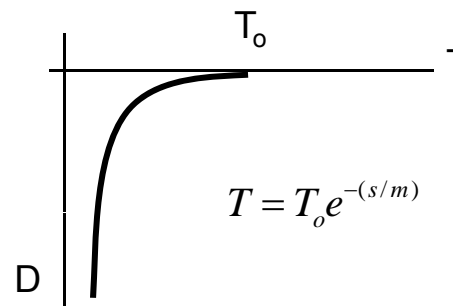
Darcy's Law:

$$q = K \frac{dh}{dx}$$

$$q(Dc) = K(Dc) \tan(\beta)$$

$$Q = Tc \tan(\beta)$$

Transmissivity vs Depth



s - soil moisture deficit

Problems:

Precipitation Uniform

Instantaneous Redistribution







## Last Year in Review

$$AR = Q = Tc \tan(\beta)$$

$$\frac{AR}{c \tan(\beta)} = T_o e^{-(s/m)} \quad a = \frac{A}{c} \quad \text{flow accumulation}$$

$$\frac{Ra}{T_o \tan(\beta)} = e^{-(s/m)}$$

$$s = -m \left[ \ln\left(\frac{R}{T_o}\right) + \ln\left(\frac{a}{\tan(\beta)}\right) \right] \quad \ln\left(\frac{a}{\tan(\beta)}\right) - \text{Topographic Index}$$

$$\bar{s} = -m \left[ \ln\left(\frac{R}{T_o}\right) + \lambda \right] \quad \lambda - \text{Average Topographic Index}$$

Simple Algebra

$$s_i - \bar{s} = -m \left[ \ln\left(\frac{a}{\tan(\beta)}\right)_i - \lambda \right]$$

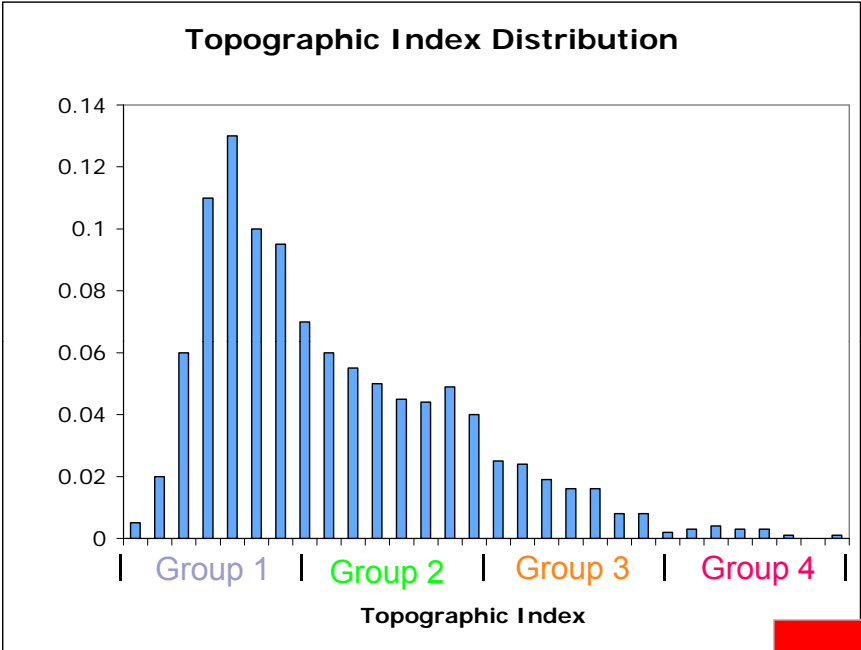
known from DEM

**TOPMODEL**  
**Linear Redistribution Equation** ☺





# Last Year in Review

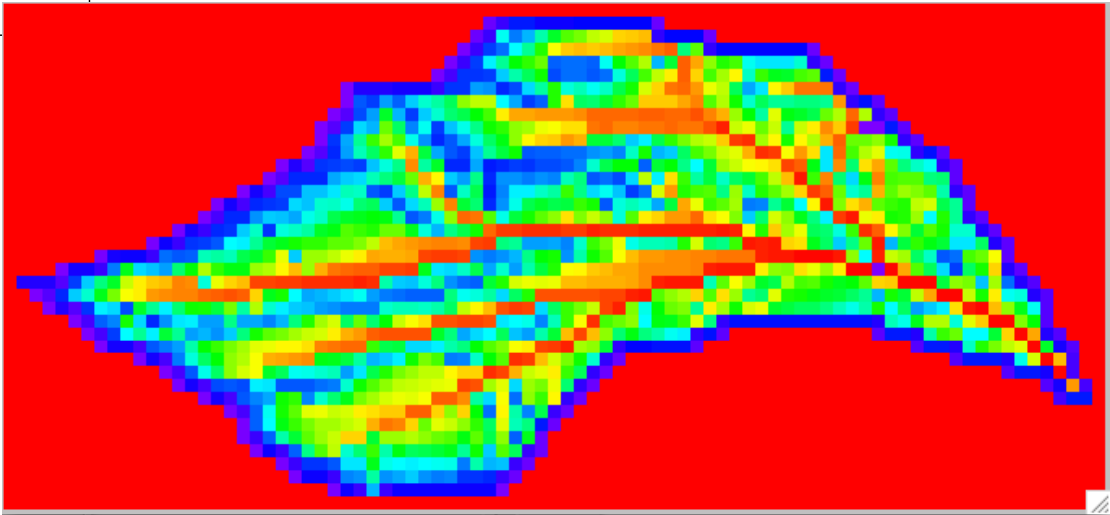


Hydrologic Similarity via Slope Position

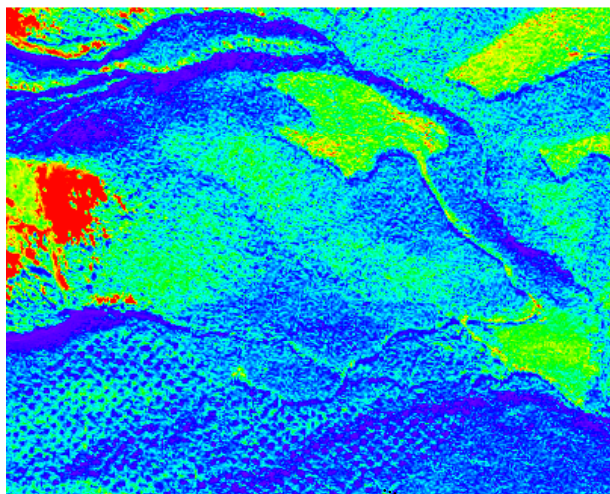
Moisture Distribution via Average Storage

Measurable Parameters [ $T_o$ ,  $m$ ,  $\ln(a/\tan(\beta))$ ]

Topographic Index  
Marmot Creek

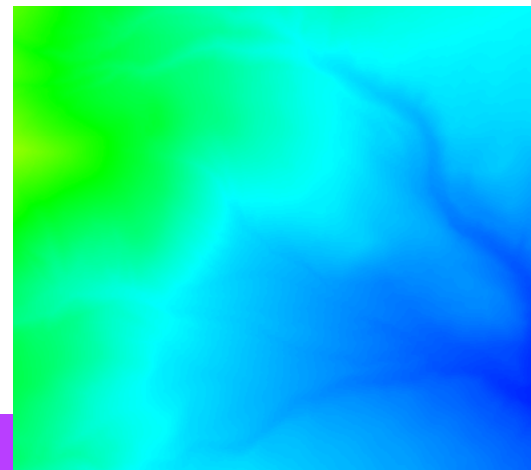


# Need For Hydrologic Similarity

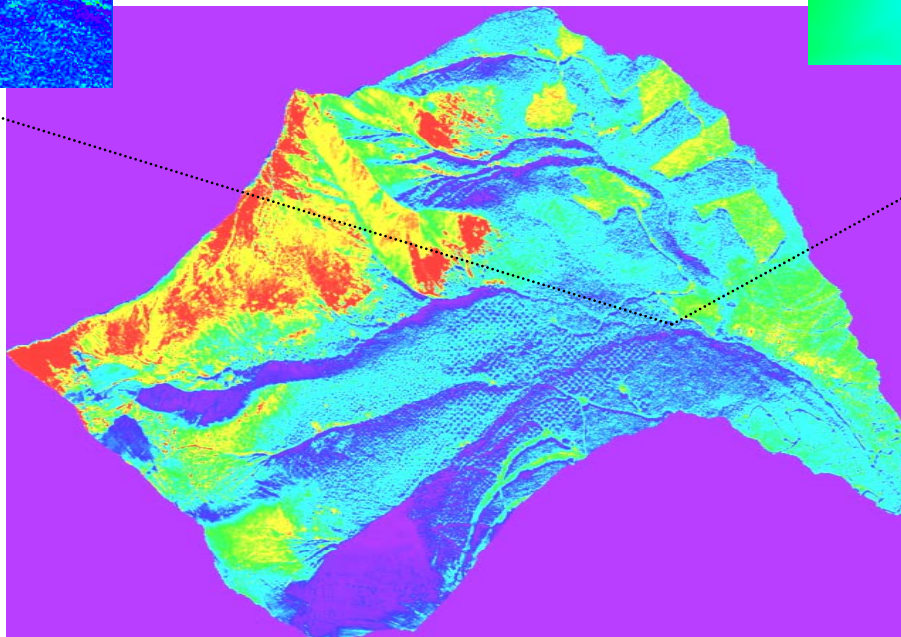


4 m Land Cover

How will we summarize and utilize all the spatial data?



1 m Topography





# Topographic Index

- ◆ **What is topographic index?**

$\ln(a/\tan\beta)$

$a$  is specific contributing area

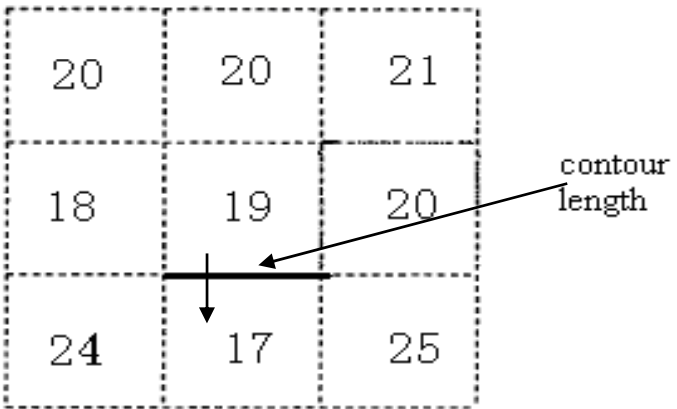
$\tan\beta$  is ground surface slope

- ◆ **Flow routing algorithms**

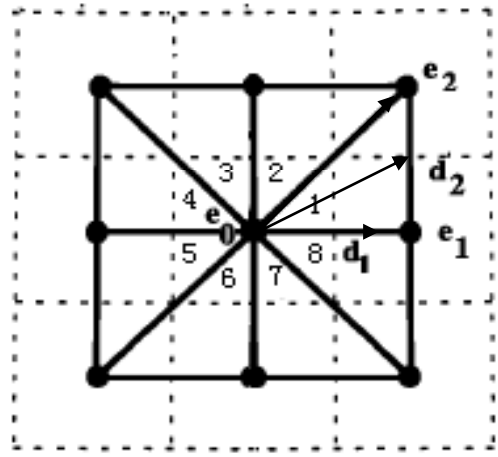
Single flow direction algorithm(D8)

Biflow direction algorithm ( $D^\infty$ )

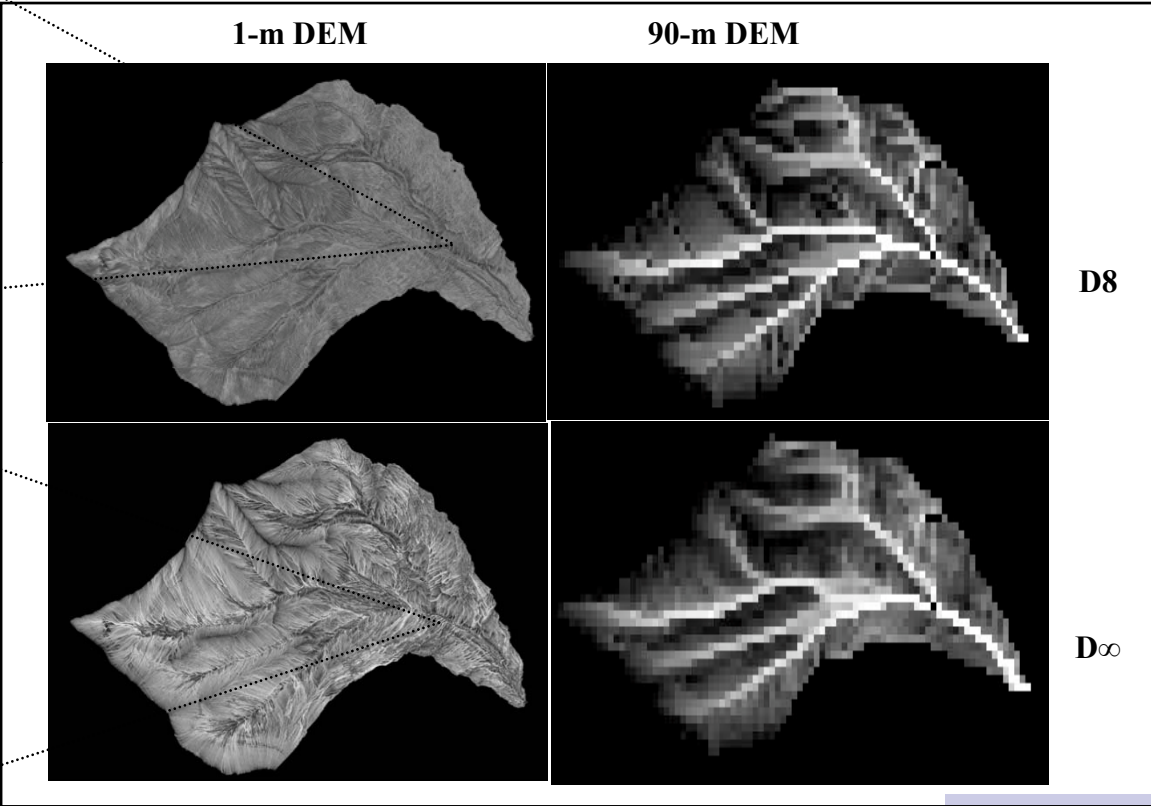
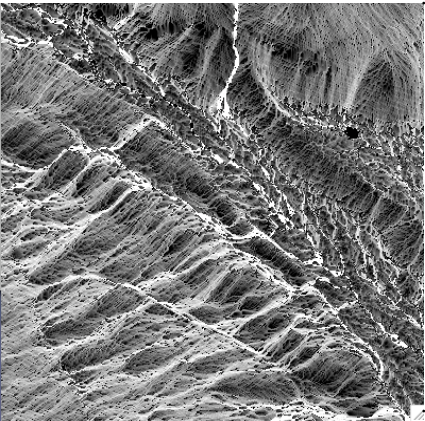
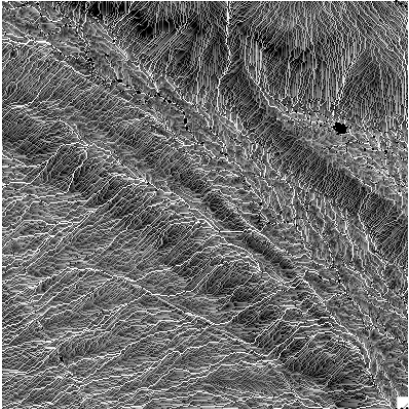
**D8**



**$D^\infty$**

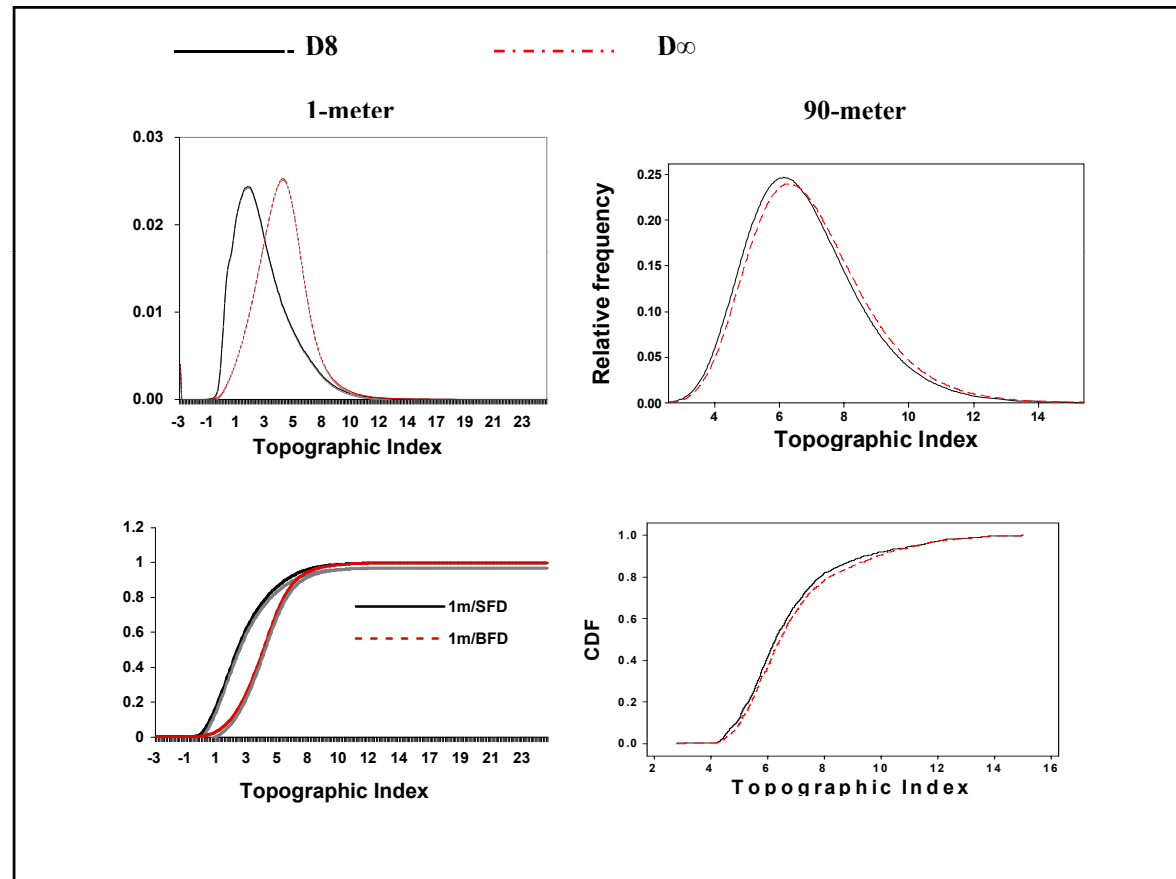


# Topographic Index Results



# Topographic Index Results

- Frequency Distributions of  $\ln(a/\tan\beta)$
- No big difference between 90-meter resolution DEM
- The discrepancy of 1-meter resolution is obvious





## Topographic Index Results

Grid Size	Algorithm	Variable	Mean
90-meter	D8	$\ln(a/\tan \beta)$	6.77
	$D^\infty$	$\ln(a/\tan \beta)$	6.95
1-meter	D8	$\ln(a/\tan \beta)$	3.05
	$D^\infty$	$\ln(a/\tan \beta)$	4.50

◆ D8 →  $D^\infty$   $\ln(a/\tan \beta)$  ↑

◆ 1-meter → 90-meter DEM  $\ln(a/\tan \beta)$  ↑

# CLASS & Resistance Calcs

## ◆ What is CLASS?

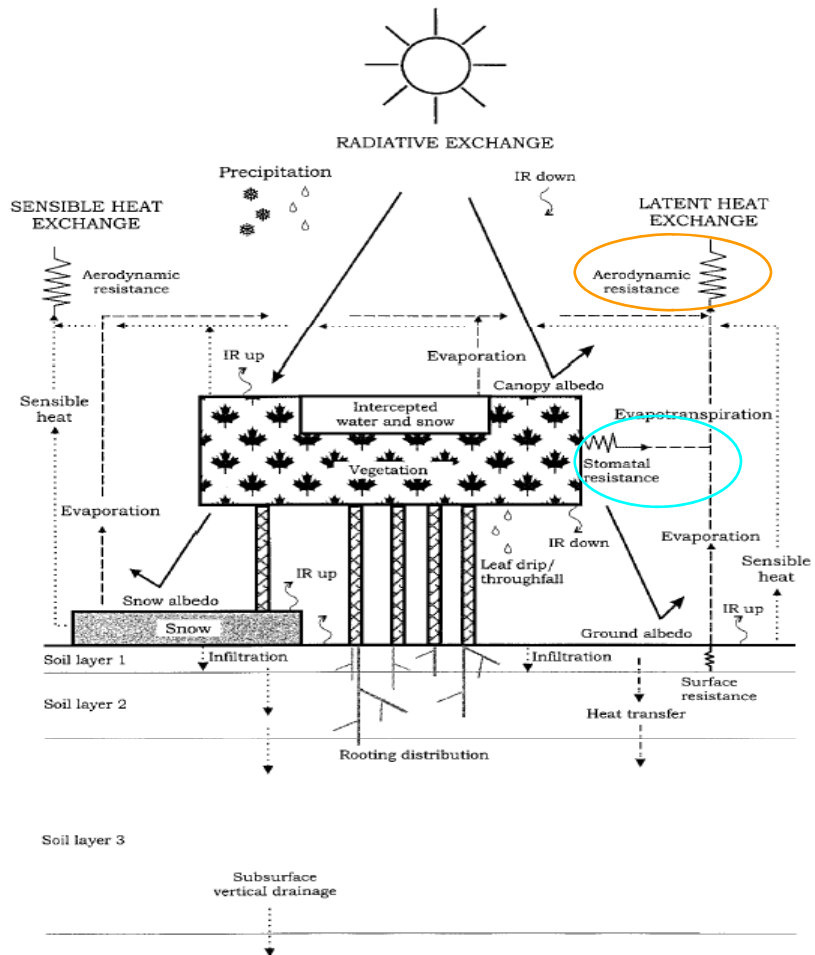
Canadian Land Surface Scheme is first developed in 1987 at the Meteorological Service

## ◆ What is the aim of CLASS

To simulate the energy and water balances of vegetation, snow and soil

## ◆ What is the objective of CLASS simulation in this study?

To simulate stomatal resistance and aerodynamic resistance



Schematic diagram of CLASS (from Versegny, 2000)





# Stomatal Resistance

## ◆ Stomatal resistance is estimated as:

$$r_{c,i} = f_1(T_{air}) f_2(\Delta e) f_3(\psi_s) r_{c,u,i}$$

$$r_{c,u} = r_{s,min} \kappa e / \ln[(K \downarrow + K \downarrow / 2 / \kappa e) / (K \downarrow \exp(-\kappa e \Lambda) + K \downarrow / 2 / \kappa e)]$$

$$\begin{aligned} f_1(T_{air}) &= 1.0 \\ &= 250.0 \\ &= 1.0 / [1.0 - (5.0 - T_{air}) \times 0.1] \\ &= 1.0 / [1.0 - (T_{air} - 40.0) \times 0.1] \end{aligned}$$

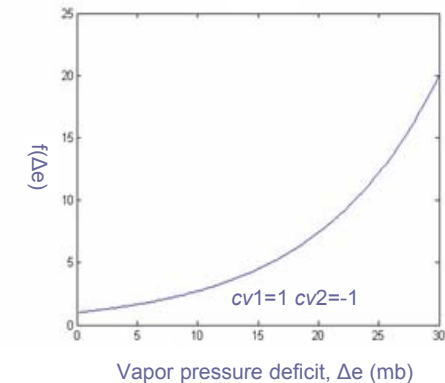
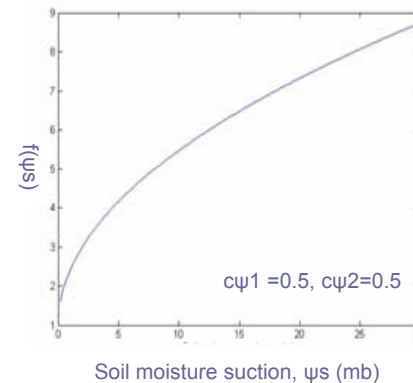
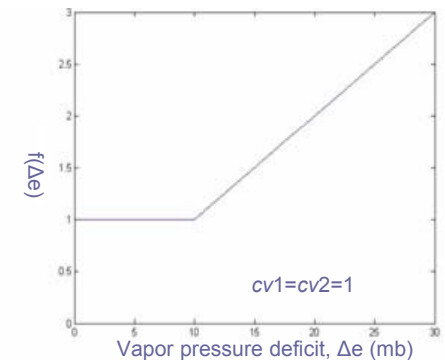
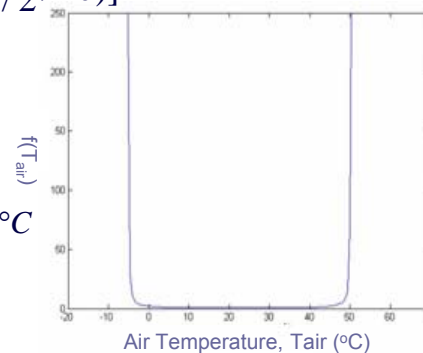
$$\begin{aligned} 5^\circ\text{C} &\leq T_{air} \leq 40^\circ\text{C} \\ T_{air} &\geq 50^\circ\text{C}, \text{ or } T_{air} \leq -5^\circ\text{C} \\ -5^\circ\text{C} &< T_{air} < 5^\circ\text{C} \\ 40^\circ\text{C} &< T_{air} < 50^\circ\text{C} \end{aligned}$$

$$\begin{aligned} f_2(\Delta e) &= (\Delta e / 10.0)^{cv2 / cv1} \\ &= 1 / \exp(-cv1 \Delta e / 10.0) \end{aligned}$$

$$\begin{aligned} cv2 &> 0 \\ cv2 &\leq 0 \end{aligned}$$

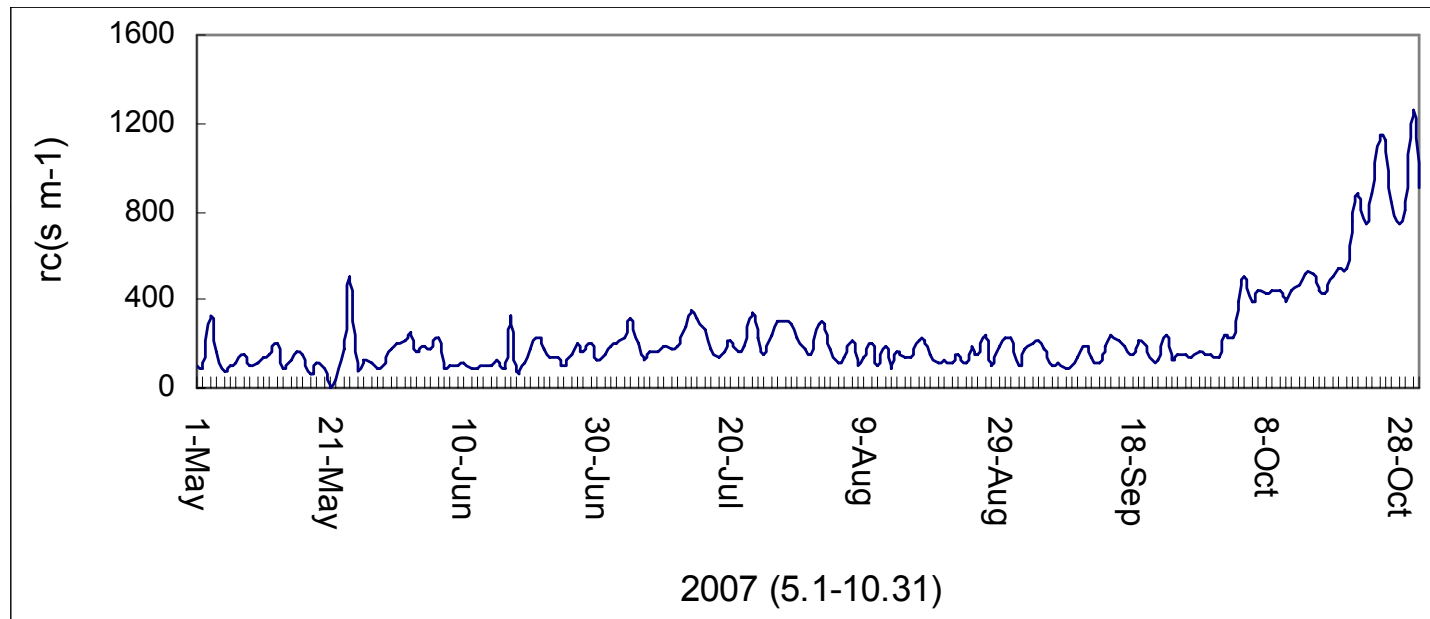
$$\begin{aligned} f_3(\psi_s) &= 1 + (\psi_s / c\psi1)^{c\psi2} \\ &= 1 \end{aligned}$$

For PET Calculation





- ◆ The estimate of stomatal resistance



- ◆ Period from 9:00 a.m. to 5:30 p.m.
- ◆ Low in May-September, High in October
- ◆ Mean value is 267 s m<sup>-1</sup>



# Aerodynamic Resistance

- ◆ **What is aerodynamic resistance ( $r_a$ )?**

The resistance encountered by fluxes of water vapor or heat or momentum along the path of transfer, which is from the source to a given reference air level above

- ◆ **How does CLASS determine aerodynamic resistance?**

$$r_a = \frac{1}{C_D v_a}$$

Where  $C_D$  is the surface drag coefficient, and  $v_a$  is the wind speed



## Aerodynamic Resistance (cont'd)

- ◆ How to determine drag coefficient for heat and water vapour fluxes  $C_{D,E}$ ?

$$C_{D,E} = \left[ \frac{k}{\ln\left(\frac{z_m - z_d}{z_{0,E}}\right)} \right] \left[ \frac{k}{\ln\left(\frac{z_m - z_d}{z_{0,M}}\right)} \right] \Phi_E \Phi_M$$

Where  $z_m$  (m) is the reference height;  $z_d$  (m) is the zero-plane displacement;  $k=0.04$  is von Karman's constant;  $z_{0,M}$  (m) is the roughness length for momentum transfer;  $z_{0,E}$  (m) is the roughness length for heat or vapor pressure transfer;  $\Phi_M$  and  $\Phi_E$  are stability correction factors.



## Aerodynamic Resistance (cont'd)

- ◆ Is there another method to determine aerodynamic resistance?

$$r_{ah0} = \frac{\left[ \ln\left(\frac{z_m - z_d}{z_0}\right) \right]^2}{k^2 v_a}$$

- ◆ Is this method available under all the conditions?

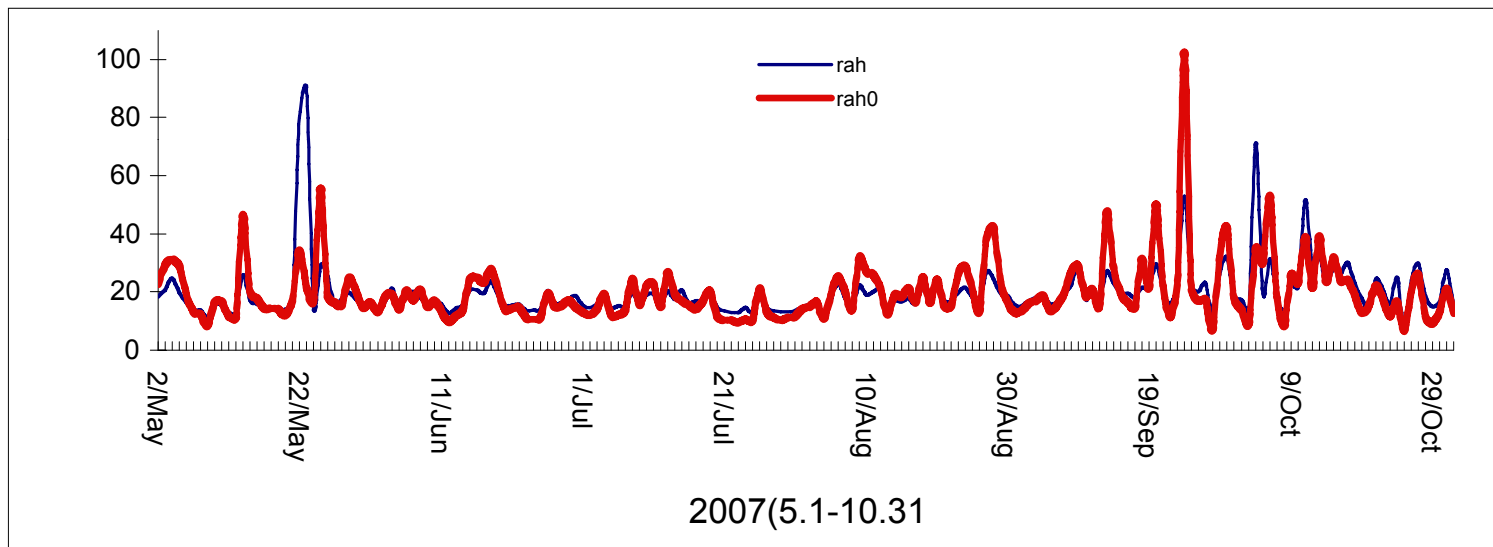
It is only available under neutral condition.

$\Phi_M = \Phi_E = 1$  and  $z_{0,E} = z_{0,M}$ , these two methods are the same



## Aerodynamic Resistance Results

- ◆ The estimate of aerodynamic resistance



- ◆ Period from 9:00 a.m. to 5:30 p.m.
- ◆ Mean value of  $r_{ah}$  is  $20.52 \text{ s m}^{-1}$
- ◆ Mean value of  $r_{ah0}$  is  $19.83 \text{ s m}^{-1}$       Note:  $r_c$  values average  $270 \text{ s m}^{-1}$



- ◆ **What is potential evapotranspiration (PET)?**

PET expresses as the amount of water that could evaporate and transpire from a vegetated landscape without restrictions other than the atmospheric demand.

- ◆ **How can we estimate PET?**

Lysimeters

Eddy correlation

Theoretical or empirical equation

Multiplying standard pan evaporation data by a coefficient

- ◆ **Which method do we use in this study?**

Penman-Monteith Equation



# Potential Evapotranspiration

## ◆ What is the Penman-Monteith Method?

$$PET = \frac{\Delta \cdot (K + L) + \rho_a \cdot c_a \cdot C_{at} \cdot e_a^* \cdot (1 - W_a)}{\rho_w \cdot \lambda_v \cdot [\Delta + \gamma \cdot (1 + C_{at} / C_{can})]}$$

$e_a^* = 0.611 \cdot \exp\left(\frac{17.3 \cdot T}{T + 237.3}\right)$

$\Delta = \frac{2508.3}{(T+237.3)^2} \cdot \exp\left(\frac{17.3 \cdot T}{T + 237.3}\right)$

$\lambda_v = 2.50 - 2.36 \times 10^{-3} \cdot T$

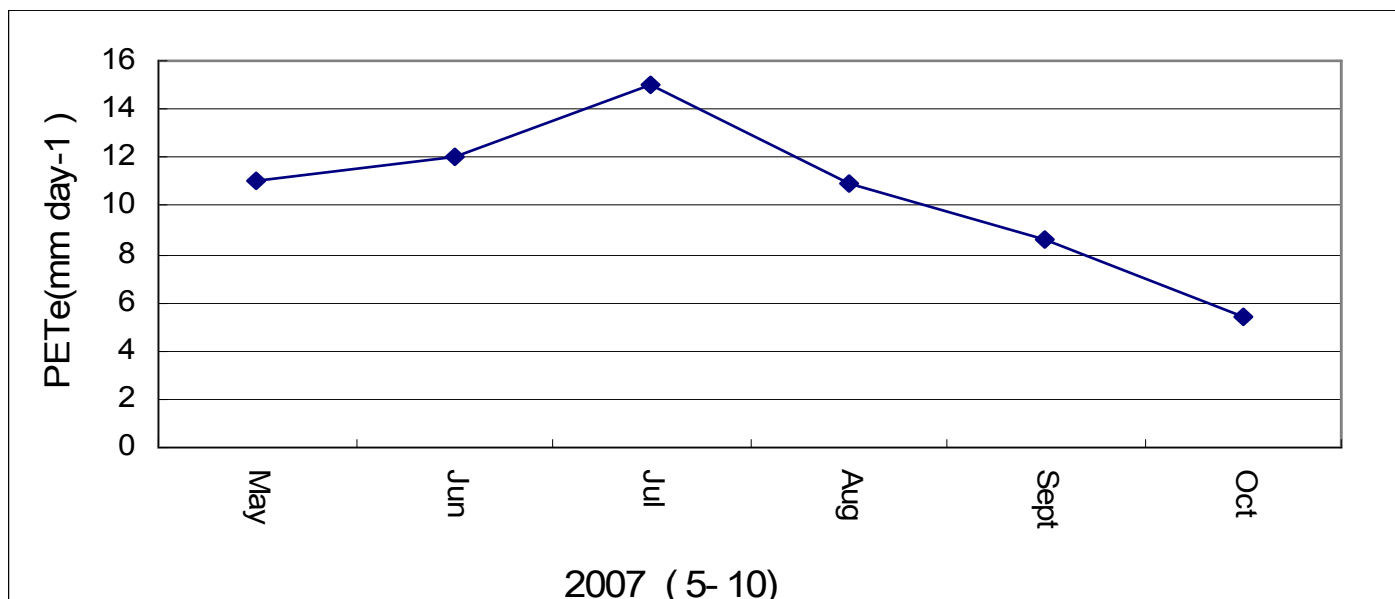
$\gamma \equiv \frac{c_a \cdot P}{0.622 \cdot \lambda_v}$

$C_{can} = f_s \cdot LAI \cdot C_{leaf}$





- ◆ PET seasonal trends in 2007

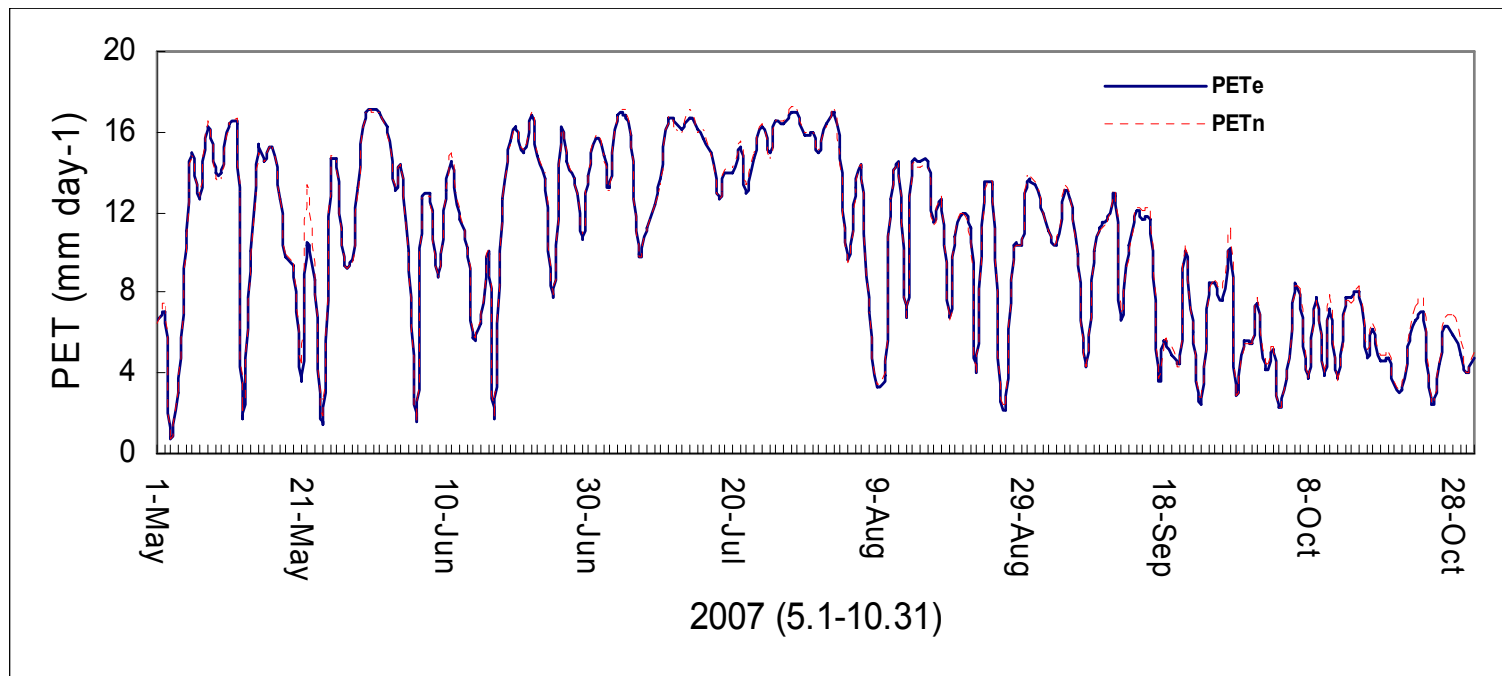


- ◆ Mean value is 10.52 mm day<sup>-1</sup>
- ◆ Highest in July, lowest in October



## PET Results (cont'd)

- ◆ Aerodynamic resistance effect on PET



- ◆ Sign test shows that PET<sub>e</sub> and PET<sub>n</sub> is not significant different
- ◆ PET is insensitive to aerodynamic resistance
- ◆ Mean value of  $r_{ah}$  is 20.52 s m<sup>-1</sup> , the mean value of  $r_c$  is 267s m<sup>-1</sup>



- ◆ **What is TOPMODEL?**

A physically-based model based on the concept of variable source area

- ◆ **Why we used TOPMODEL in this study?**

It can be used to predict streamflow, overland and subsurface flow, and soil moisture deficit with less parameters

- ◆ **What is the objective for this study?**

To simulate runoff in the Marmot Creek catchment using TOPMODEL



## TOPMODEL (cont'd)

- ◆ How does TOPMODEL determine streamflow?

$$q_{streamflow} = \frac{\sum a_i p + \sum a_i |s_i|}{A} + T_0 e^{-\lambda} e^{-\frac{\bar{s}}{m}}$$

where  $a_i$  is specific area,  $p$  is precipitation,  $s_i$  is local soil moisture deficit,  $A$  is watershed area,  $T_0$  is saturated transmissivity,  $\lambda$  is the mean  $\ln(a/\tan\beta)$  for the catchment,  $\bar{s}$  is catchment-average saturation deficit, and  $m$  is parameter

- ◆ How does TOPMODEL determine soil moisture deficit ( $s_i$ )?

$$s_i = \bar{s} + m \left[ \lambda - \ln \left( \frac{a}{\tan \beta_x} \right) \right]$$



## TOPMODEL (cont'd)

### ◆ How many parameters involved in TOPMODEL?

Parameter	Name in theory	Description	Range
<b>SZM[m]</b>	m	Transmissivity decay parameter	0.005-0.06
<b>T0 [m<sup>2</sup>h<sup>-1</sup>]</b>	ln(T <sub>0</sub> )	Effective lateral saturated transmissivity	0.1-8
<b>TD [m h<sup>-1</sup>]</b>	t <sub>d</sub>	Unsaturated zone time delay	0.1-500
<b>CHV [m h<sup>-1</sup>]</b>	CHV	Channel velocity	100-10000
<b>RV [m<sup>-2</sup> h<sup>-1</sup>]</b>	RV	Routing velocity	100-10000
<b>SRMAX [m]</b>	SRmax	Maximum allowable root zone storage	0.005-0.3
<b>SR0[m]</b>	SR0	Initial root zone deficit	0.0-0.3



## TOPMODEL (cont'd)

- ◆ **How to calibrate the parameters?**

Genetic Algorithm

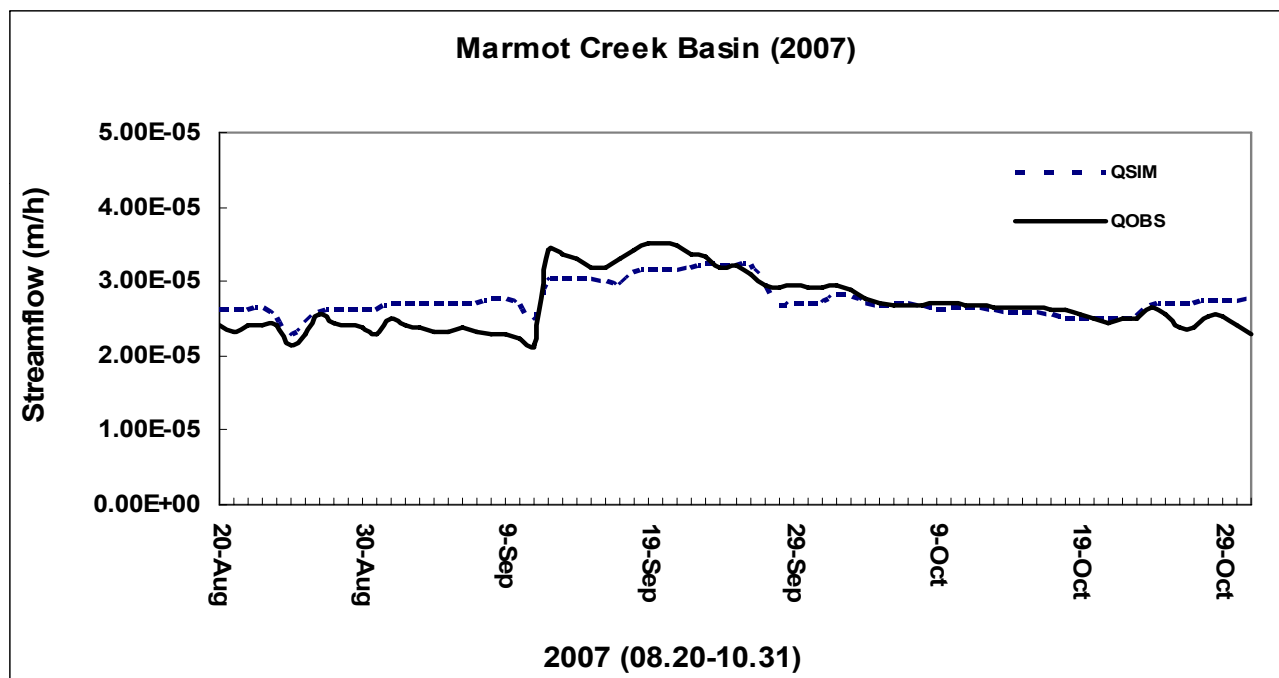
- ◆ **How to evaluate the model efficiency?**

$$EFF = \frac{\sum(QOBS_i - QOBS_m)^2 - \sum(QOBS_i - QSIM_i)^2}{\sum(QOBS_i - QOBS_m)^2}$$

Where  $QOBS_i$  is the observed streamflow,  $QSIM_i$  is the simulated streamflow, and  $QOBS_m$  is the mean of the observed streamflow.



- ◆ The result of calibration



- ◆ Period of 20th August to 31st October in 2007
- ◆ The model efficiency value is 0.611



## TOPMODEL Results (cont'd)

- ◆ Calibration results for 1-meter and 90-meter resolution DEM, using D8 and D<sup>∞</sup>.

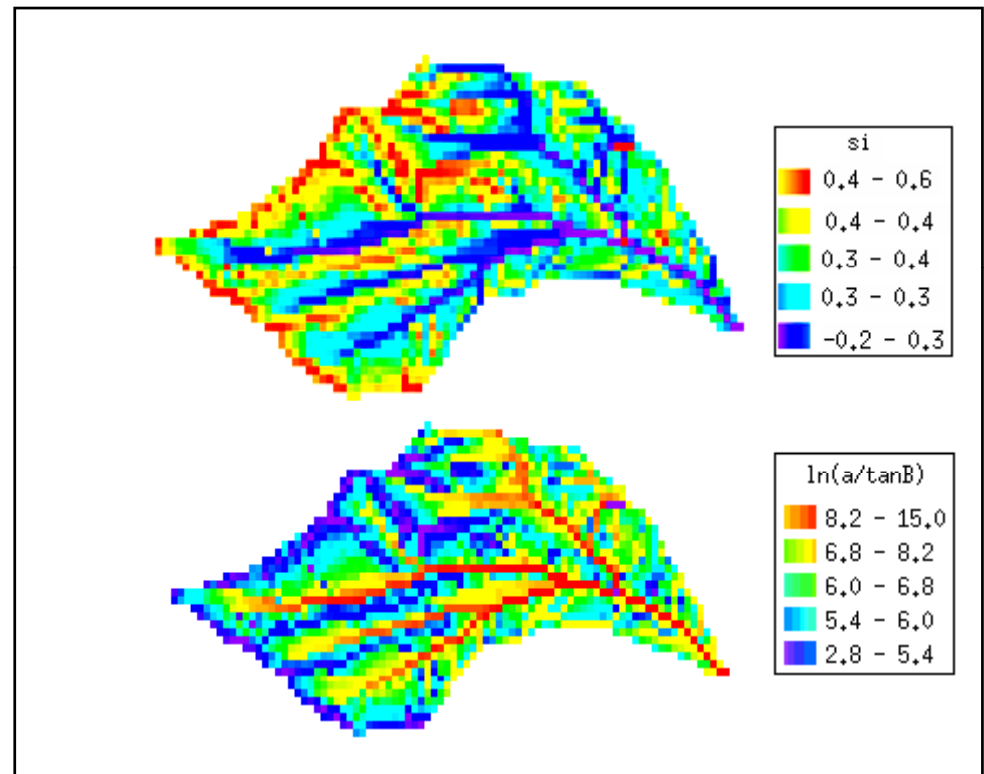
	Flow	SZM (m)	T0 (m <sup>2</sup> h <sup>-1</sup> )	SR0 (m)	SRMAX (m)	CHV (m h <sup>-1</sup> )	RV (m <sup>2</sup> h <sup>-1</sup> )	Td (mh <sup>-1</sup> )	EFF	QSUB (%)
<b>1</b>	<b>D8</b>	0.046	0.29	0.0015	0.0059	371.53	449.81	382.24	0.611	84.91
	<b>D<sup>∞</sup></b>	0.046	0.29	0.0015	0.0059	371.53	449.81	382.24	0.583	86.4
<b>90</b>	<b>D8</b>	0.060	1.18	0.00006	0.0060	1725.00	114.84	261.85	0.656	52.55
	<b>D<sup>∞</sup></b>	0.060	1.18	0.00006	0.0060	1725.00	114.84	261.85	0.565	53.96





## TOPMODEL Results (cont'd)

- ◆ The relation between soil moisture deficit( $si$ ) and topographic index [ $\ln(a/\tan\beta)$ ] on August 28th in 2007.
- ◆ Large value of  $\ln(a/\tan\beta)$  indicates the locations within a watershed most likely to be saturated and produce overland flow





# Land Cover Classes Vs Topographic Index

- ◆ What are the procedures of land cover classification?

Images processing

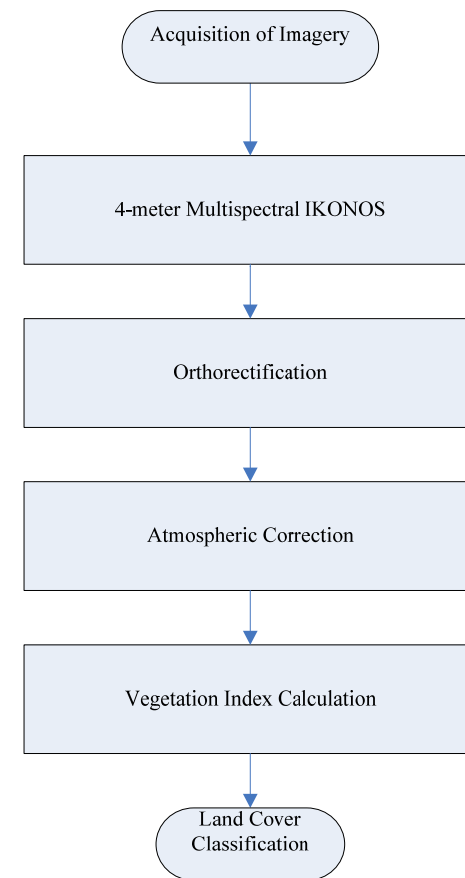
Calculation of vegetation index

Land cover classification

- ◆ Which Normalized Difference Vegetation Index (NDVI) is used in this study?

$$NDVI = \frac{\rho_{nir} - \rho_{red}}{\rho_{nir} + \rho_{red}} \quad -1.0 \leq NDVI \leq 1.0$$

Where  $\rho_{nir}$  and  $\rho_{red}$  represent reflectance at the red and near-infrared (NIR) wavelengths, respectively.



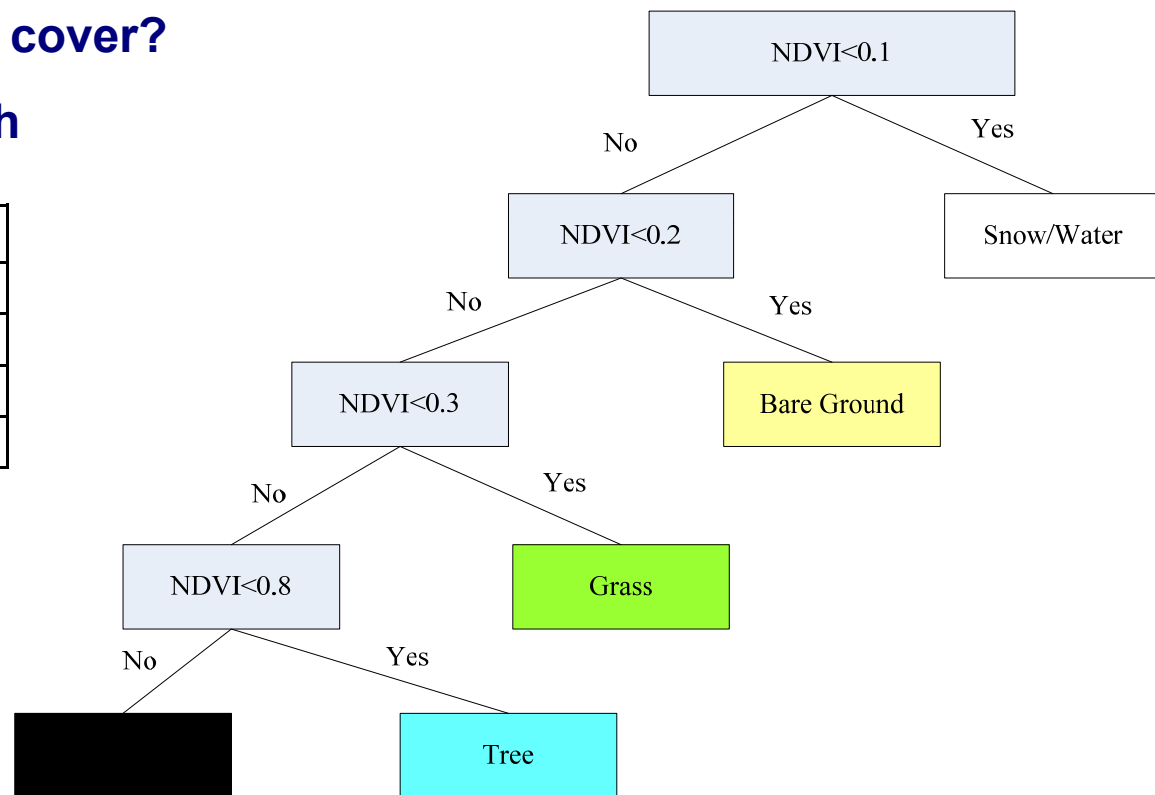


# Land Cover Classes Vs Topographic Index

◆ How can we classify the land cover?

Using Decision tree approach

Land Cover Classes	NDVI Rang
Snow or Water	$NDVI < 0.1$
Bare Ground	$0.1 \leq NDVI < 0.2$
Grass	$0.2 \leq NDVI < 0.3$
Tree	$0.3 \leq NDVI < 0.8$

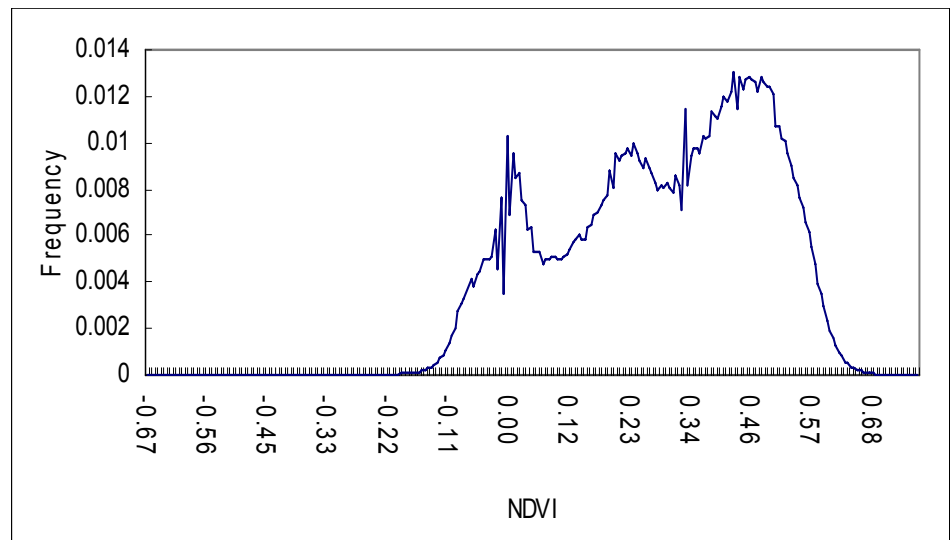
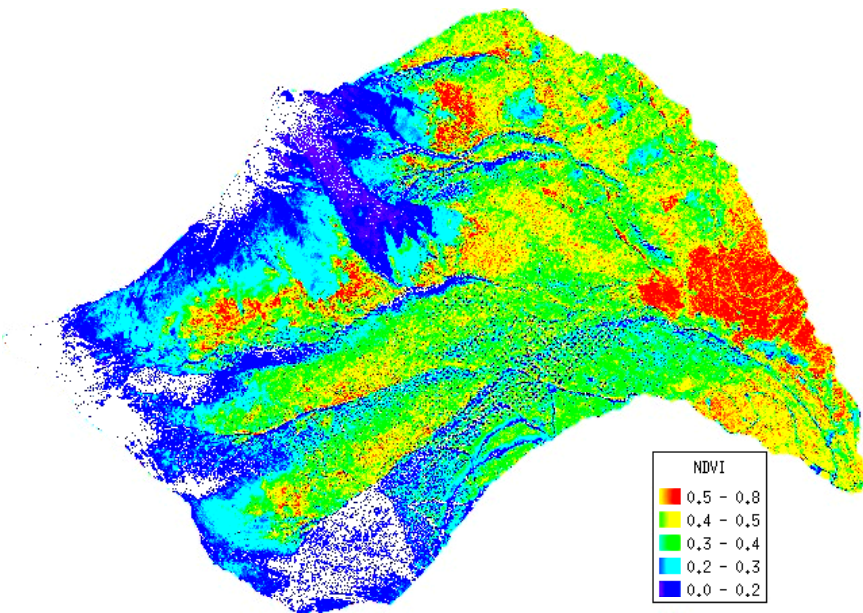




# NDVI Calculation Results

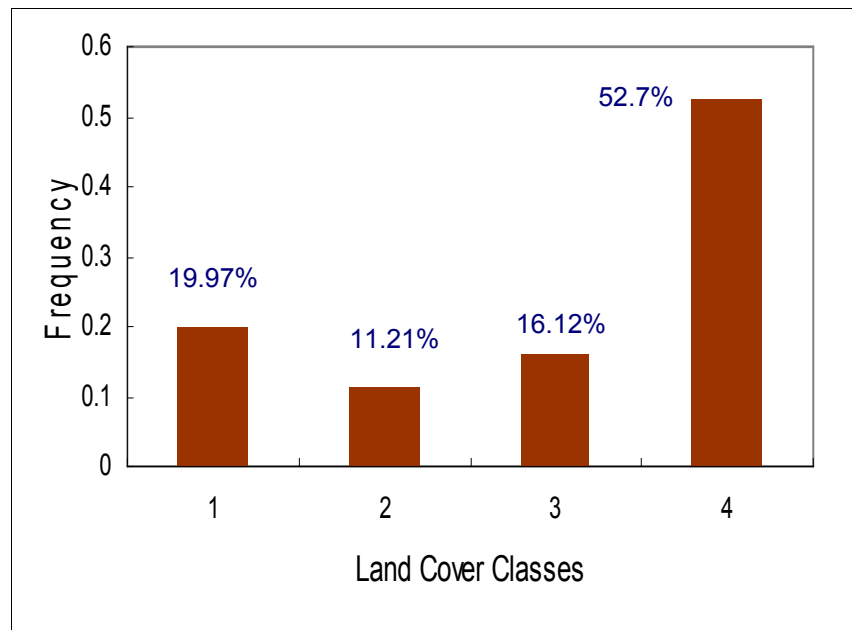
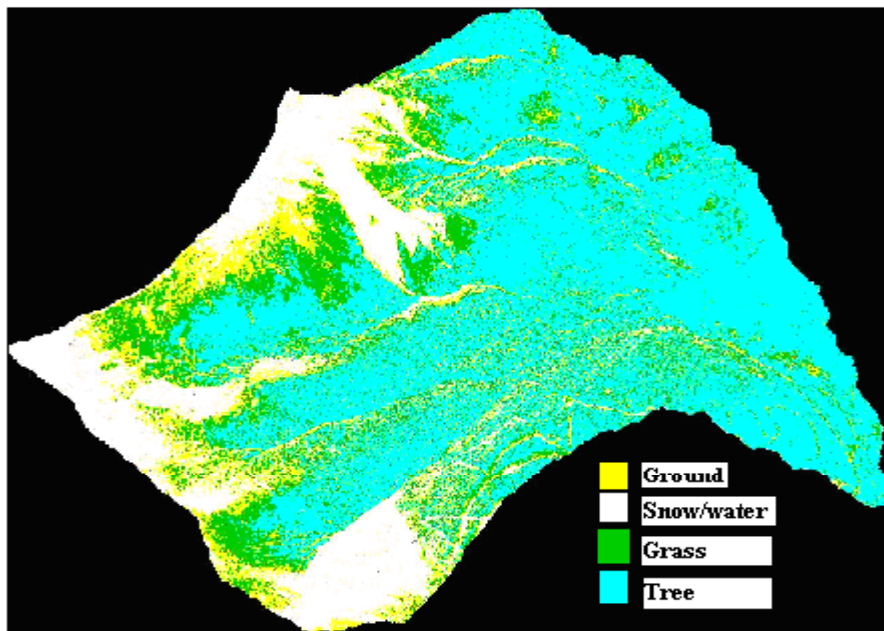
- ◆ **The results of NDVI calculation**

NDVI in this study area varies from -0.67 to 0.77, mean value is 0.29, in October 18<sup>th</sup>, 2003.



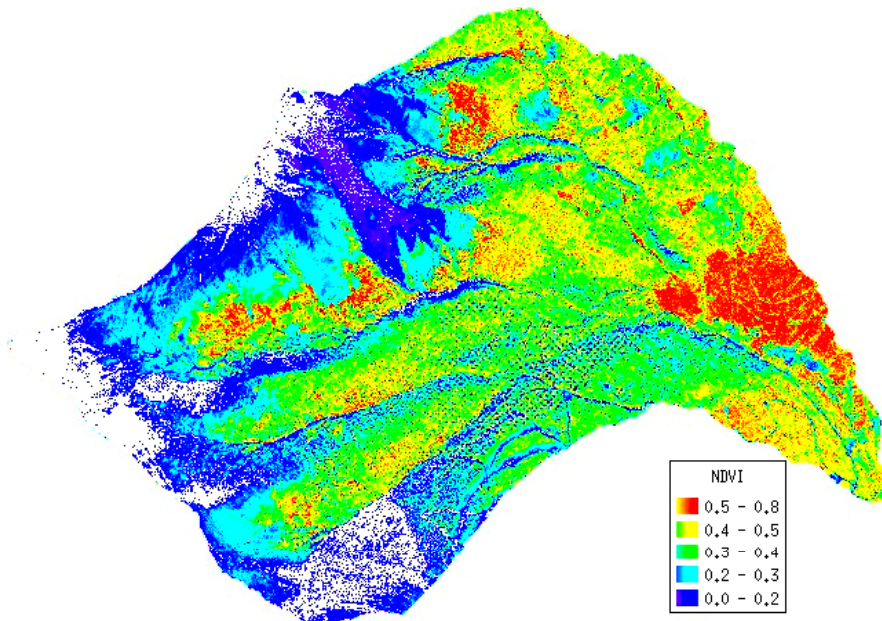


◆ The results of land cover classification

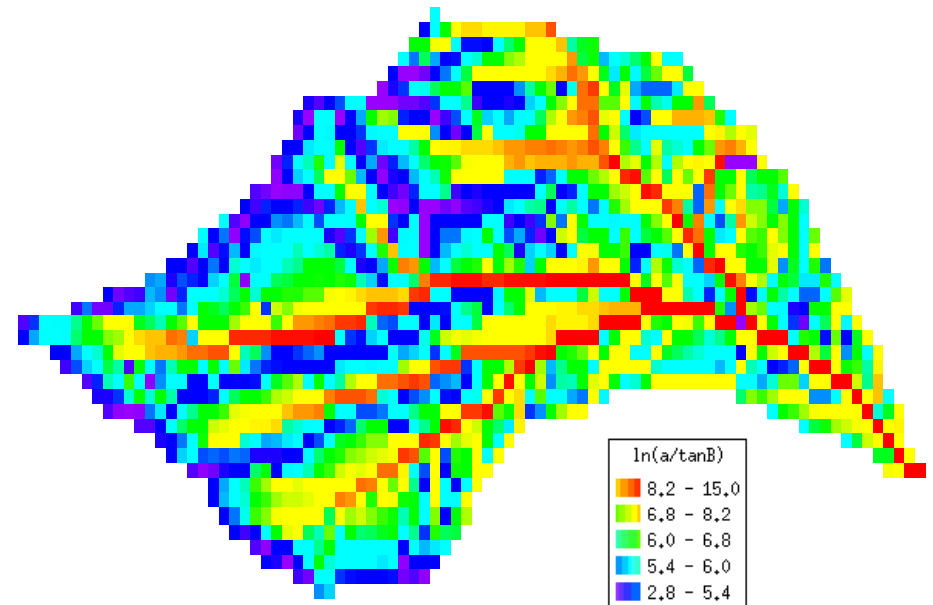




# Developing New Measures of Similarity



NDVI - Land Cover



Topographic Index - Terrain



## Future work

### Future Work

#### ◆ Theme II - Parameterization

Potential evapotranspiration estimates

Topographic Index Calculation - scaling behavior

Develop Similarity Estimates - terrain, topo, elevation, aspect

#### ◆ Theme III - Prediction

Implement TOPMODEL redistribution within CLASS

Prediction of finer scale topography