# Modelling the spatial variability of the snowcover and the surface energy fluxes at TVC

S. Endrizzi, P. Marsh, and S. Pohl







## **Purpose of Study**

- Study the effects of topography and land cover on snow cover variability and surface energy balance
- Improve the understanding of the processes causing spatial variability of turbulent fluxes
- Improve the parameterization of these processes at the small scale

#### **Snow cover variability at TVC**

- Variability <u>during accumulation</u>:
  - Wind redistribution
  - Snow interception by vegetation
- Variability <u>during melting</u>:
  - Effect of topography on radiation
  - Effect of wind variability on turbulent fluxes
  - Effect of vegetation (shrubs) on the surface energy balance

#### To understand processes...

- Measurements of <u>surface energy fluxes</u> at 2 stations (TUP and TTS) - including turbulent fluxes - during and after snow
- Several Landsat <u>SCA maps</u> during snowmelt
- <u>Maps of turbulent fluxes</u> derived from data from NRCC Twin Otter aircraft carrying EC sensor

#### To implement the parameterizations...

- Used the **<u>GEOtop</u>** distributed hydrological model
  - It is <u>grid based</u>, so it represents well the effect of topography, normally it is run at a smaller scale (10-500m than LSMs)
  - Fully describes the surface energy balance (radiation and turbulent fluxes)
  - Discretizes snow and soil with several layers (the user can choose according to the )
  - Solves heat equation with phase change in the soil and snow column coupled with water flow equation (Richards' equation in the soil)
  - Fully 3D Richards' equation solved with iterative methods -> calculates runoff

## **GEOtop**



- State variables:
  - Temperature
  - Water content
  - Ice content
- Non linearity problems

soil water equation  $c_{rit}(\psi)\frac{\partial\psi(\theta)}{\partial t} = \frac{\partial}{\partial z}\left(-k(\psi) + k(\psi)\frac{\partial\psi(\theta)}{\partial z}\right) + s(\psi)$ 

heat equation with phase change  $C(\theta_w, \theta_{ice}) \frac{\partial T}{\partial t} - L\rho_{ice} \frac{\partial \theta_{ice}}{\partial t} = \frac{\partial}{\partial z} \left( k_T (\theta_w, \theta_{ice}) \frac{\partial T}{\partial z} \right)$ 

#### **GEOtop for the application in the arctic**

- Coupled with <u>PBSM</u> and with <u>MicroMet</u> to represent a wind field including topographic effects and to describe <u>snow transport</u> by wind (of paramount importance in TVC)
- Implemented a <u>dual layer scheme</u> to describe the effects of shrubs on the surface energy balance, in particular the turbulent fluxes



The parameterization of the "undercanopy" resistance is critical because of uncertainty and high sensitivity

- <u>Measurements</u> of turbulent fluxes of sensible and latent heat have been performed at TUP and TTS towers during spring since 2003, during and after snowmelt
- The models has been applied to:
  - understand how shrub tundra affect turbulent exchange in different conditions
  - test to ability of the model to reproduce the fluxes



# **Canopy fraction variable in time** (to account for <u>gradual</u> shrub burial)

$$f_c = 1 - \left(\frac{D_{sn}}{D_{snB}}\right)^q \le 1$$

 $f_c$  = canopy fraction

 $D_{snB}$  = threshold snow depth above which all the shrubs are considered buried has been defined.

# *q* = exponent (>1) which takes into account that the probability of shrub emergence increases more than linearly the snow depth decreases

In the snow melt period, when the snow depth  $(D_{sn})$  becomes smaller than  $D_{snB}$ , the shrubs are considered to gradually emerge from the snow cover



#### **DURING MELTING**



#### TTS Shrub Tundra



TUP Open Tundra

- Despite the reduction in the wind speed, the shrubs <u>increase the</u> <u>contribution of sensible heat to the surface energy balance</u> and decrease the heat losses - good agreement observations - results
- This of of great importance in snowmelt acceleration is shrubs tundra
- The results depend much on the parameterization of the undercanopy resistance. Stable stratifications may occur in the canopy reducing the turbulence exchange.



#### AFTER MELTING





TTS Shrub Tundra



- Underestimation at TTS good estimation at TUP
- Higher sensible fluxes at TTS, higher turbulence
- On the other hand, latent heat fluxes are higher at TUP (vegetation is still non transpiring)

#### AFTER MELTING

LE







TUP Open Tundra



Data from NRCC **Twin Otter** aircraft carrying EC sensor during the 1999 melting season

Mauder et al. (2008) -> **2D maps** at <u>3 km</u> resolution (footprint model + wavelet method)

Maps for 4 days (<u>**27 May</u> - 1 Jun -**3 Jun - 8 Jun)</u>

#### Snow water equivalent (mm) 27 May 1999 - SCA approx 70%



#### Land cover:

Tundra (white) - Shrubs (yellow) - Forest (green) - Lakes (cyan) - No data (red)



#### From aircraft

Modelled





#### 8th June 1999

#### From aircraft

Η

#### Modelled





#### 8th June 1999

#### From aircraft

LE

#### Modelled





#### Discussion

- Reproduced main spatial patterns of sensible flux during snowmelt (higher in snow free areas and over shrubs, lower in snow covered open tundra)
- After snowmelt: measurements at TUP and TTS suggest negative correlation between LE and H; the aircraft data the contrary
- This arises question. Is the model good enough in predicting the fluxes 2D? Are aircraft data really representative of the surface?

### **Snow spatial variability**

Snow water equivalent - SCA

#### SWE (mm) - 23 May 1999 - model 20 m



#### SCA - 23 May 1999 - SPOT



## Conclusion

- The model allows interpreting the measurements of turbulent fluxes separating the fluxes in canopy and undercanopy components
- The model performs reasonably well and the temporal patterns of the fluxes are in generally well reproduced, during and after snowmelt
- The parameterization of the undercanopy resistance should be improved
- The spatial patterns of turbulent fluxes have to be further investigated, given disagreements between model and aircraft results

# Thank you very much